

**GEOLOGIC MAP OF THE FARMINGTON QUADRANGLE,
SALT LAKE AND DAVIS COUNTIES, UTAH**
by
Mike Lowe², Stefan M. Kirby¹, and Kimm M. Harty²
2018

ADJOINING 7.5' QUADRANGLE NAMES



DESCRIPTION OF MAP UNITS

QUATERNARY

Alluvial deposits

Qal, **Younger stream alluvium** (upper Holocene) – Clast-supported, moderately to well-sorted pebble and cobble gravel; gravely sand, and silty sand deposited in modern channels and floodplains; clasts subangular to rounded; deposited where fluvial processes are currently or episodically active; locally includes minor alluvial-fan, colluvial, and terrace deposits; thickness generally less than 6 meters (20 ft).

Qal₁, **Older stream alluvium** (middle Holocene to upper Pleistocene) – Clast-supported, moderately sorted pebble and cobble gravel; gravely sand, and silty sand deposited along incised floodplains 1 to 3 meters (3–10 ft) above modern stream level; clasts subangular to rounded; mapped along Haight Creek and adjacent drainages in the northern part of the quadrangle where fluvial processes are generally no longer active; thickness generally less than 6 meters (20 ft).

Qalp, **Stream alluvium related to the Provo shoreline and regressive phase of Lake Bonneville** (upper Pleistocene) – Clast-supported, moderately to well-sorted pebble and cobble gravel; gravely sand, and silty sand; deposited along inactive floodplains more than 3 meters (10 ft) above modern stream level; mapped where fluvial processes are generally no longer active; exposed thickness less than 6 meters (20 ft).

Qal_f, **Younger alluvial-fan deposits** (upper Holocene) – Mixture of gravel and sand deposited by streams, and dissection deepened by debris flows; forms fans, locally with distinct levees and channels, at mouths of mountain-front canyons; exposed thickness less than 6 meters (20 ft).

Qal_f, **Older alluvial-fan deposits** (middle Holocene to upper Pleistocene) – Mixture of gravel and sand deposited by streams, and dissection deepened by debris flows; forms fans that are slightly incised by modern stream channels; exposed thickness less than 6 meters (20 ft).

Qalp, **Alluvial-fan deposits related to the Provo shoreline and regressive phase of Lake Bonneville** (upper Pleistocene) – Mixture of gravel and sand deposited by streams, and dissection deepened by debris flows; forms fans graded approximately to the Provo level of Lake Bonneville that are incised by modern stream channels; exposed thickness less than 6 meters (20 ft).

Artificial deposits

Qld, **Fill and disturbed land** (historical) – Land disturbed and excavated through aggregate (sand and gravel) operations and construction of Interstate highways, highway interchanges, and Farmington Bay dikes.

Colluvial deposits

Qc, **Colluvium** (Holocene to middle Pleistocene) – Weakly to non-layered, variably sorted, matrix- to clast-supported silt, sand, clay, and minor gravel of local origin; deposits formed mostly by creep and slope wash; includes hill-slope deposits of angular cobble to pebble-sized clasts of Farmington Canyon Complex float; thickness probably less than 3 meters (10 ft) in most areas.

Wetland and marsh deposits

Qsm, **Wetland and marsh deposits** (Holocene) – Wet, fine-grained, organic-rich sediment associated with springs, wetlands, ponds and seeps; thickness probably less than 1 meter (3 ft) in most areas.

Mass-movement deposits

Qmf, **Debris-flow deposits** (upper Holocene) – Matrix- to clast-supported cobble and boulder gravel, with variable amounts of sand, silt, and clay matrix; surfaces variably rubbly and commonly have levees and channels; thickness probably less than 9 meters (30 ft).

Qmf₁, **Older debris-flow deposits** (middle to lower Holocene) – Matrix- to clast-supported cobble and boulder gravel, with variable amounts of sand, silt, and clay matrix; surfaces variably rubbly and commonly have levees and channels; includes multiple events graded to various levels above modern channels; unit grades into alluvial fans at mouths of canyons; thickness probably less than 9 meters (30 ft).

Qml, **Liquefaction-induced landslide deposits** (Holocene) – Mixture of silt, fine sand, and minor gravel redeposited in flow slides and lateral spreads (the northern part of the Farmington Siding landslide complex) as a result of liquefaction during large earthquakes; deposits display landslide-related lineaments, scarps, and hummocky topography; disrupted bedding and sand-filled cracks (injection features) are present in the deposits in the subsurface (Hyland and Lowe, 1998; Hartly and Lowe, 2003); thickness generally less than 22 meters (70 ft). Most recent movement between 2700 and 4500 cal yr B.P.

Qml, **Older liquefaction-induced landslide deposits** (lower Holocene to upper Pleistocene) – Mixture of silt, fine sand, and minor gravel redeposited in flow slides and lateral spreads (the southern part of the Farmington Siding landslide complex) as a result of liquefaction during large earthquakes; deposits display landslide-related lineaments, scarps, and hummocky topography that are more subdued than similar features found in the northern part of the landslide complex (Hyland and Lowe, 1998; Hartly and Lowe, 2003); thickness less than 22 meters (70 ft). Most recent movement between 11,500 and 13,000 cal yr B.P.

Qms, **Landslide deposits** (upper Holocene) – Unsorted, unstratified mixtures of gravel, sand, silt, and clay redeposited by slides, slumps, and flows; deposits display distinctly hummocky topography and fresh scarps, and are currently or have been recently active; thickness generally less than 15 meters (50 ft).

Qms, **Older landslide deposits** (middle to lower Holocene) – Unsorted, unstratified mixtures of mostly sand, silt, and clay redeposited by single to multiple slides, slumps, and flows; deposits display hummocky topography but lack fresh scarps and are mostly inactive; thickness generally less than 15 meters (50 ft).

Lacustrine deposits

Qli, **Lacustrine silt and clay deposits** (Holocene) – Silt and clay with minor sand deposited in mud flats and exposed by fluctuating Great Salt Lake levels; may contain gypsum, halite, and other salts; thickness typically less than 3 meters (10 ft).

Qlf, **Lacustrine fine-grained deposits** (upper Pleistocene) – Intervals of mixed fine-grained sediment, clay to silt, and intervals of rhythmically interbedded clay to medium sand deposited in low-energy, generally offshore environments at elevations below the Provo shoreline; thickness typically less than 6 meters (20 ft).

Qlsp, **Lacustrine sand-bearing deposits related to the Provo shoreline and regressive phase of Lake Bonneville** (upper Pleistocene) – Moderately to well-sorted gravely sand, interlayered with some silt and sand; deposited and reworked in higher energy environments along the Provo and regressive shorelines near the mountain front; mapped at elevations below Provo shoreline; thickness typically less than 6 meters (20 ft).

Qlspb, **Lacustrine sand-bearing deposits related to transgressive and regressive phases of Lake Bonneville** (upper Pleistocene) – Moderately to well-sorted gravely sand, interlayered with some silt and sand; deposited and reworked in higher energy environments along the Provo and regressive shorelines near the mountain front; mapped at elevations below Provo shoreline; thickness typically less than 6 meters (20 ft).

Qlspb, **Lacustrine gravel-bearing deposits related to transgressive and regressive phases of Lake Bonneville** (upper Pleistocene) – Sand and clast-supported, moderately sorted pebble to cobble gravel, mapped below the Provo shoreline; gravels contain rounded to subrounded clasts, and some subangular clasts derived from reworking of mass-movement and alluvial-fan deposits; deposited in higher energy environments along shorelines; thickness generally less than 9 meters (30 ft).

Qlspb, **Lacustrine gravel-bearing deposits related to the Bonneville shoreline and transgressive phase of Lake Bonneville** (upper Pleistocene) – Clast-supported, moderately to well-sorted pebble to cobble gravel, with some silt to sand in interflow areas and away from mountain front; gravels contain rounded to subrounded clasts, and some subangular clasts derived from reworking of mass-movement and alluvial-fan deposits; deposited in higher energy environments along shorelines during Lake Bonneville transgression; thickness generally less than 9 meters (30 ft).

Mixed-environment deposits

Qlas, **Mixed alluvial and marsh deposits** (Holocene) – Predominantly fine-grained sediment (sand, silt, and clay) deposited by low-gradient streams and in marshes; total thickness typically less than 6 meters (20 ft).

Q, **Quaternary unconsolidated basin fill** (Holocene to Pleistocene) – Unconsolidated mixture of lacustrine and alluvial clay, silt, sand, gravel, marl, and tuffaceous layers; shown only on cross sections; up to 150 meters (500 ft) thick.

Qb, **Quaternary basin fill** (Pleistocene) – Weakly consolidated mixture of lacustrine and alluvial clay, silt, sand, gravel, marl, and tuffaceous layers; shown only on cross sections; up to 400 meters (1300 ft) thick.

Tb, **Tertiary basin fill** (Tertiary) – Weakly to strongly consolidated mixture of conglomerate, sandstone, mudstone, tuffaceous sandstone, tuff, and lacustrine limestone; shown only on cross sections; up to 2400 meters (8000 ft) thick.

PALEOPROTEROZOIC

Xfsg, **Farmington Canyon Complex quartz-rich gneiss** (Paleoproterozoic) – Medium- to light-gray, moderately foliated and layered, quartz-feldspathic gneiss. Unit is characterized by quartz, feldspar, hornblende, and biotite grains, with garnet porphyroblasts; unit also contains zones of migmatite, pegmatitic dikes, amphibolite layers, biotite-rich schist, and syenite. Unit forms rubble-strewn cliffs and steep slopes along Farmington Canyon and to the south along the mountain front. Age of metamorphism is ~1700 Ma (Nelson and others, 2002; Mueller and others, 2011; Nelson and others, 2011).

Xfsm, **Farmington Canyon Complex pegmatitic gneiss** (Paleoproterozoic) – Light-gray, weakly to moderately foliated, pegmatitic gneiss. Unit consists of quartz-feldspathic pegmatitic gneiss with minor biotite mica and secondary chlorite alteration. Unit may contain relatively unfoliated pegmatite bands and pods, and zones of moderately foliated interlayered pegmatite and quartz-rich gneiss. Unit is inset in quartz-rich gneiss (Xfsg) exposures south of Farmington Canyon. Age of metamorphism is ~1700 Ma (Nelson and others, 2002; Mueller and others, 2011; Nelson and others, 2011).

Xfc, **Farmington Canyon Complex, undifferentiated** (Paleoproterozoic) – Medium- to light-gray, moderately foliated and layered, quartz-feldspathic gneiss and quartz-feldspathic pegmatite gneiss; shown only on cross sections.

REFERENCES

Arnott, T., and Stephens, D., 1990, Hydrologic characteristics of the Great Salt Lake, Utah—1847-1986: U.S. Geological Survey Water-Supply Paper 2332, 32 p., scale 1:25,000.

Brook Ramsey, C., 2009, Bayesian analysis of radiocarbon dates: Radiocarbon, v. 51, no. 1, p. 337–360.

Bryant, B., 1988, Geologic map of the Farmington Canyon Complex, Wasatch Mountains, Utah: U.S. Geological Survey Professional Paper 1476, 54 p., 1 plate, scale 1:50,000.

Curry, D.R., and James, S.R., 1982, Paleoenvironments of the northeastern Great Basin and the Basin Rim region—a review of geological and biological evidence, in Madsen, D.B., and O'Connor, J.F., editors, Man and environment in the Great Basin: Society for American Archaeology Papers, no. 2, p. 27–52.

Godsey, H.S., Curry, D.R., and Chan, M.A., 2005, New evidence for an extended occupation of the Provo shoreline and implications for regional climate change, Pleistocene Lake Bonneville, Utah: Quaternary Research, v. 63, p. 212–223.

Godsey, H.S., Oviatt, C.G., Miller, D.M., and Chan, M.A., 2011, Stratigraphy and chronology of offshore to nearshore deposits associated with the Provo shoreline, Pleistocene Lake Bonneville, Utah: Palaeogeography, Palaeoclimatology, Palaeoecology, v. 310, no. 3–4, p. 442–450.

Harty, K.M., and Lowe, M., 2003, Geologic evaluation and hazard potential of liquefaction-induced landslides along the Wasatch Front, Utah: Utah Geological Survey Special Study 104, 40 p., 16 plates, scale 1:24,000.

Hyland, M.D., and Lowe, M., 1998, Characteristics, timing, and hazard potential of liquefaction-induced landsliding in the Farmington Siding landslide complex, Davis County, Utah: Utah Geological Survey Special Study 95, 38 p.

Mabey, D.R., 1992, Subsurface geology along the Wasatch Front: U.S. Geological Survey Professional Paper 1500-C, 16 p.

McGee, D., Quade, J., Edwards, R.L., Brockner, W.S., Cheng, H., Reiners, P.W., and Evenson, P., 2012, Lacustrine cave carbonates—novel archives of paleohydrologic change in the Bonneville Basin (Utah, USA): Earth and Planetary Science Letters, v. 351–352, p. 182–194.

McKean, A.P., and Hyland, M.D., 2013, Interim geologic map of the Baileys Lake quadrangle, Salt Lake and Davis Counties: Utah Geological Survey Open-File Report 624, 18 p., 1 plate, scale 1:24,000.

McNeil, B.R., and Smith, R.B., 1992, Upper crustal structure of the northern Wasatch Front, Utah, from seismic reflection and gravity data: Utah Geological Survey Contract Report 92-7, 62 p.

Miller, D.M., 2016, The Provo shoreline of Lake Bonneville, in Oviatt, C.G., and Shroder, J.F., Jr., editors, Lake Bonneville—a scientific update: Amsterdam, Netherlands, Elsevier, Developments in Earth Surface Processes, v. 20, chapter 7, p. 127–144.

Miller, D.M., Oviatt, C.G., Dadash, S.L., and McGeehin, J.P., 2005, Late Holocene highstands of Great Salt Lake at Locomotive Springs, Utah: Geological Society of America Abstracts with Programs, v. 37, no. 7, p. 335.

Mueller, P.A., Wooden, J.L., Mogk, D.W., and Foster, D.A., 2011, Paleoproterozoic evolution of the Farmington zone—implications for terrane accretion in southwestern Laurentia: Lithosphere, v. 3, no. 6, p. 401–408.

Murchison, S.B., 1989, Fluctuation history of Great Salt Lake, Utah, during the last 13,000 years: Salt Lake City: University of Utah, Ph.D. dissertation, 137 p.

Murchison, S.B., and Mulvey, W.E., 2000, Late Pleistocene and Holocene shoreline stratigraphy on Antelope Island, in King, J.K., and Willis, G.C., editors, The geology of Antelope Island: Utah Geological Survey Miscellaneous Publication 00-1, p. 78–83.

Nelson, A.R., and Personius, S.F., 1993, Surficial geologic map of the Weber segment, Wasatch fault zone, Weber and Davis Counties, Utah: U.S. Geological Survey Miscellaneous Investigations Series Map I-2199, 22 p., 1 plate, scale 1:50,000.

Nelson, S.T., Harris, R.A., Dorais, M.J., Heizler, M., Constenius, K.N., and Barnett, D.E., 2002, Basement complexes in the Wasatch fault, Utah, provide new limits on crustal accretion: Geology, v. 30, no. 9, p. 831–834.

Nelson, S.T., Hart, G.L., and Frost, C.D., 2011, A reassessment of Mojavia and a new Cheyenne Belt alignment in the eastern Great Basin: Geosphere, v. 7, no. 2, p. 513–527.

Oviatt, C.G., 1997, Lake Bonneville fluctuations and global climate change: Geology, v. 25, no. 2, p. 155–158.

Oviatt, C.G., 2014, The Gilbert episode in the Great Salt Lake basin, Utah: Utah Geological Survey Miscellaneous Publication 14-3, 20 p.

Oviatt, C.G., 2015, Chronology of Lake Bonneville, 30,000 to 10,000 yr B.P.: Quaternary Science Reviews, v. 110, p. 166–171. Appendix A supplementary data available online, <http://dx.doi.org/10.1016/j.quascirev.2014.12.016>.

Oviatt, C.G., Curry, D.R., and Miller, D.M., 1990, Age and paleoclimatic significance of the Stansbury shoreline of Lake Bonneville, northeastern Great Basin: Quaternary Research, v. 33, p. 291–305.

Oviatt, C.G., Miller, D.M., McGeehin, J.P., Zachary, C., and Mahan, S., 2005, The Younger Dryas phase of Great Salt Lake, Utah: Palaeogeography, Palaeoclimatology, Palaeoecology, v. 219, no. 3–4, p. 263–284.

Reimer, P.J., Bard, E., Bayliss, A., Beck, J.W., Blackwell, P.G., Brook Ramsey, C., Buck, C.E., Cheng, H., Edwards, R.L., Friedrich, M., Grootes, P.M., Guilderson, T.P., Halldason, H., Hoggins, L., Hattis, C., Heaton, T.J., Hoffmann, D.L., Hogg, A.G., Hughen, K.A., Kaiser, K.F., Kremer, B., Manning, S.W., Niu, M., Reimer, R.W., Richards, D.A., Scott, E.M., Southon, J.R., Staff, R.A., Turney, C.S.M., and van der Plicht, J., 2013, IntCal13 and Marine13 radiocarbon age calibration curves 0–50,000 years cal B.P.: Radiocarbon, v. 55, no. 4, p. 1869–1887.

Sack, D., 2005, Geologic map of the Clearfield 7.5' quadrangle, Davis County, Utah: Utah Geological Survey Miscellaneous Publication MP-05-4, 14 p., 2 plates, scale 1:24,000.

Salts, R.W., and Jachens, R.C., 1995, Gravity and basin-depth maps of the Basin and Range Province, western United States: U.S. Geological Survey Map GP-1012, scale 1:250,000.

Smith, R.B., and Bruhn, R.L., 1984, Intraplate extensional tectonics of the western U.S. Cordillera—implications on structural style from seismic-reflection data, regional tectonics and thermal-mechanical models of brittle-ductile deformation: Journal of Geophysical Research, v. 89, no. B7, p. 5733–5762.

Solomon, B.J., 2007, Surficial geologic map of the Kaysville quadrangle, Davis County, Utah: Utah Geological Survey Map 224, 2 plates, scale 1:24,000.

Van Horn, R., 1981, Geologic map of pre-Quaternary rocks of the Salt Lake City North quadrangle, Davis and Salt Lake Counties, Utah: U.S. Geological Survey Miscellaneous Investigations Series Map I-1762, 1 plate, scale 1:24,000.

Van Horn, R., and Crittenden, M.D., Jr., 1987, Map showing surficial units and bedrock geology of the Fort Douglas and parts of the Mountain Dell and Salt Lake City North quadrangles, Davis, Salt Lake, and Morgan Counties, Utah: U.S. Geological Survey Miscellaneous Investigations Series Map I-1762, 1 plate, scale 1:24,000.

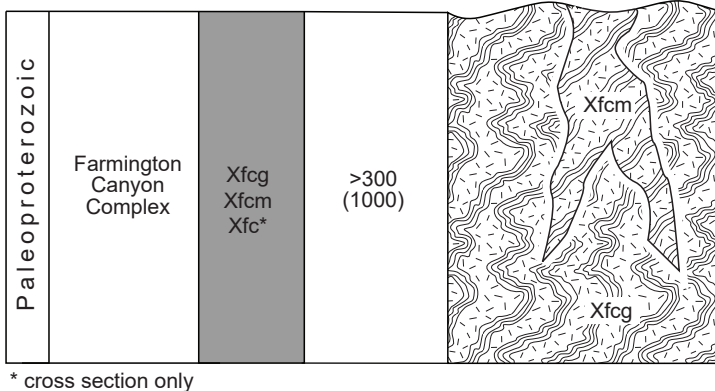
Willis, G.C., Yonkee, W.W., Doelling, H.H., and Jensen, M.E., 2010, Geology of Antelope Island State Park, Utah, in Sprinkel, D.A., Chidsey, T.C., Jr., and Anderson, F.B., editors, Geology of Utah's Parks and Monuments: Utah Geological Association Publication 28, p. 241–254.

Yonkee, W.A., 1992, Basement over relations, Sevier orogenic belt, northern Utah: Geological Society of America Bulletin, v. 104, no. 3, p. 280–302.

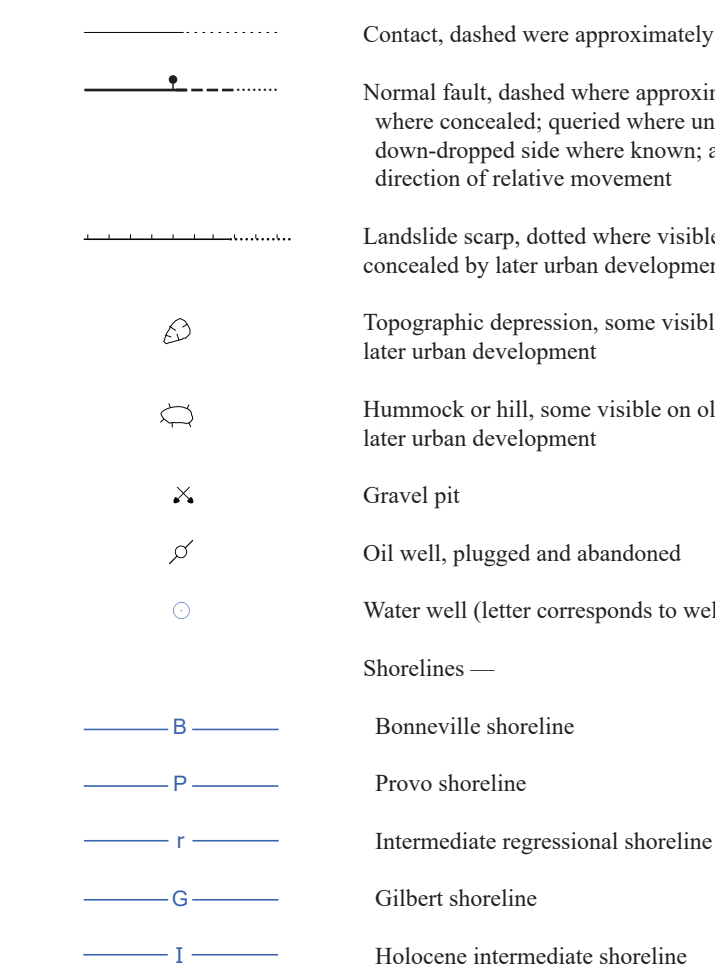
Zoback, M.L., 1983, Structure and Cenozoic tectonism along the Wasatch fault zone, Utah, in Miller, D.M., Todd, V.R., and Howard, K.A., editors, Tectonic and stratigraphic studies of the eastern Great Basin: Geological Society of America Memoir 157, p. 3–26.

LITHOLOGIC COLUMN

UNIT/FORMATION	MAP SYMBOL	THICKNESS meters (ft)	LITHOLOGY
High-Pleistocene surface deposits	Q*	variable	
Provo shoreline basin fill	Qb*	variable	
Regressive basin fill deposits	Tb*	variable	



MAP SYMBOLS



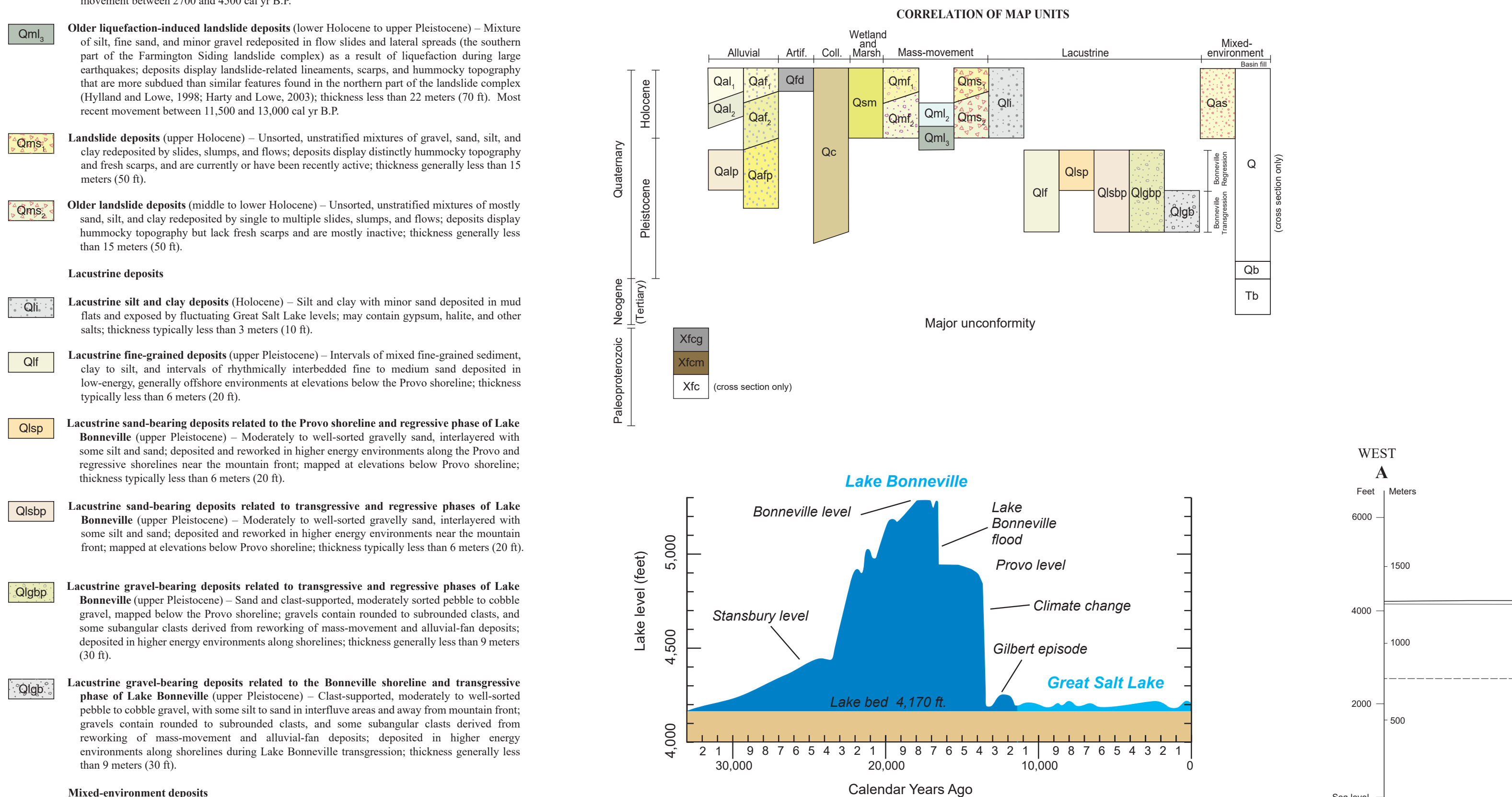
View to the southwest across the Farmington quadrangle from near the mouth of Farmington Canyon. Banded gneiss typical of the Farmington Canyon Complex is in the foreground.



Damage caused in Farmington by 1983 Bulld Canyon debris flow.



Mike Love examining exposure of Wasatch fault zone near the mouth of Davis Canyon just east of Farmington quadrangle.

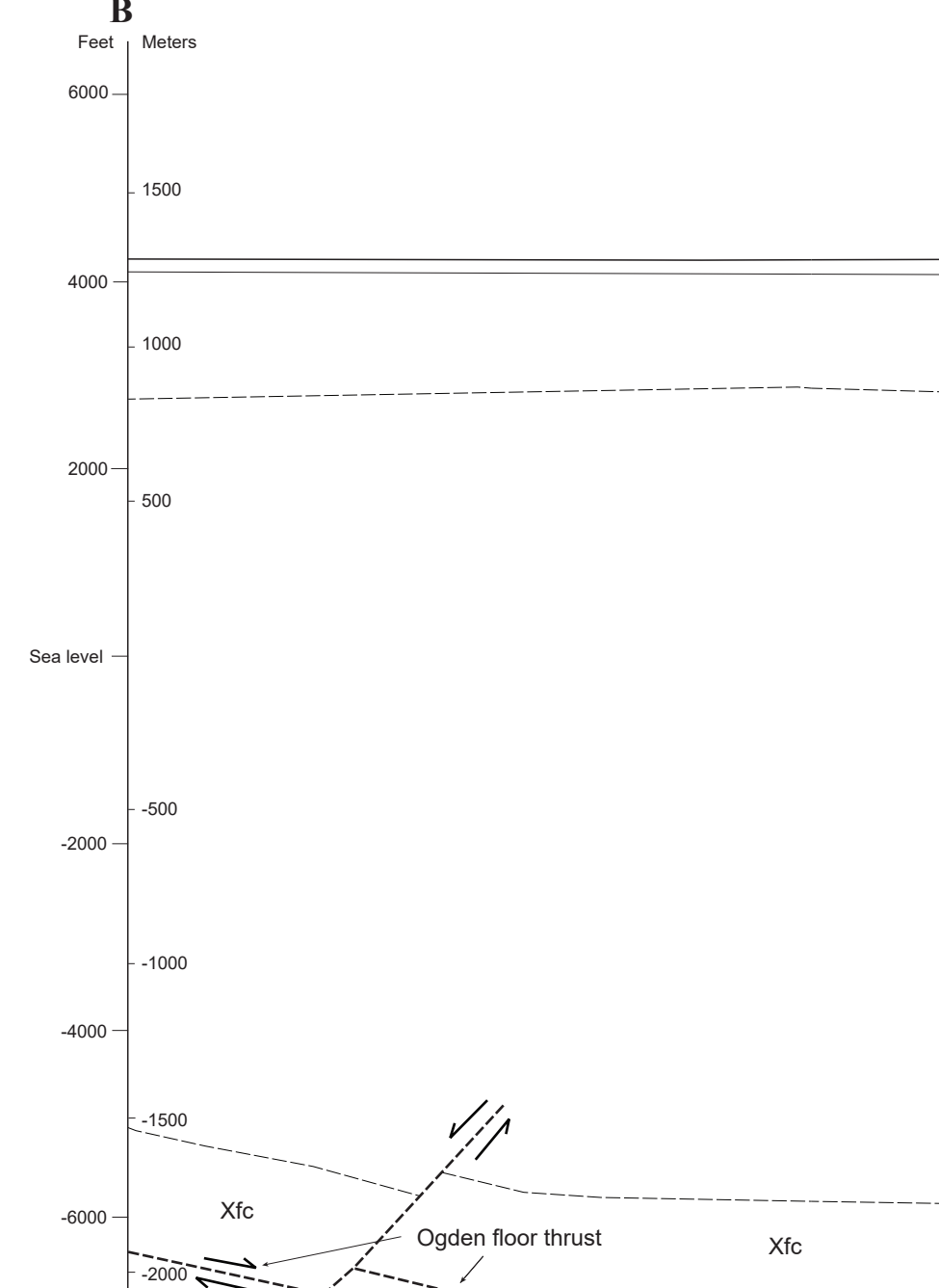


Rise and fall of Lake Bonneville between about 28 and 13.5 ka, and of Great Salt Lake (modified from Oviatt, 1997; Murchison and Mulvey, 2000; and Godsey and others, 2005). Elevations have been adjusted to remove effects of isostatic rebound. Holocene portion of hydrograph is largely schematic.

Table 1. Well location is shown on geologic map. Depth is the total depth of well in feet. WR location corresponds to the water right location in the Utah Division of Water Rights database available at <http://www.waterrights.utah.gov/wellrights/default.asp>. API number corresponds to petroleum well logs available at <http://www.oem.utah.gov/>. Location data projection: UTM 12 N AD 1927.

Well ID	Depth (ft)	Location	API or WR Location	Bottom Lithology
A	3525	111°55'47.711°W 40°57'29.605°N	43011 (0041)	Tertiary basin fill
B	985	111°56'53.31°W 40°59'59.341°N	S840 D1334 NW 13 T33N R1W	unconsolidated deposits
C	910	111°54'1.207°W 40°59'41.708°N	N100 W2579 E4 13 T33N R1W	unconsolidated deposits
D	520	111°52'36.133°W 40°52'37.861°N	S2485 W1525 NE 30 T2N R1E	unconsolidated deposits

SOUTHWEST



Cross section by Stefan M. Kirby interpreted from available seismic and gravity data (Zoback, 1983; Smith and Bruhn, 1984; Mabey, 1992; McNeil and Smith, 1992; Salts and Jachens, 1995). Structure of Ogden thrust system based in part on schematic cross sections in Yonkee (1992) and Willis and others (2010). Actual position of Ogden thrust system in the subsurface uncertain.

Table 2. Ages of major shoreline occupations of Lake Bonneville, Gilbert episode, and Great Salt Lake with shoreline elevations in the Farmington quadrangle.

Lake Cycle and Phase	Shoreline (map symbol)	Age radiocarbon years (±1σ yr B.P.)	Age calibrated years (cal yr B.P.)	Shoreline Elevation feet (meters)
Lake Bonneville				
Transgressive phase	Stansbury shorelines	22,000–20,000 ¹	26,000–24,000	Not recognized ²
	Bonneville (B)	–15,200–15,000 ³	–18,500–18,000	5180–5200 (1580–1585)
Overflowing phase	Provo (P)	–15,000–12,600 ⁴	18,000–15,000	4820–4860 (1470–1482)
Regressive phase	Regressive shorelines (r, f)	–12,600–11,500 ⁵	15,000–13,000	4380–4820 (1335–1470)
Gilbert episode	Gilbert (G)	10,000 ⁶	11,500	–4250 (1295)
Great Salt Lake				
	early Holocene highstand	9700–9400 ⁷	11,000–10,500	Not recognized
	late Holocene highstand	4200–2100 ⁸	5000–2000	4217–4221 (1285–1287)
	Historical highstand		late 1860s to early 1870s and 1986–87 ⁹	4212 (1284)

¹All calibrations made using OxCal 4.3.2 calibration and analysis software (version 4.3.2; Bronk Ramsey, 2009; using the IntCal13 calibration curve of Reimer and others, 2013), rounded to the nearest 500 years.

²Oviatt and others (1990)

³The Stansbury shoreline formed at elevations of about 4440 to 4450 feet (1350–1360 m), which are present in the quadrangle, but the shoreline was either weakly developed or poorly preserved and cannot be identified.

⁴Oviatt (2015), Miller (2016), and references therein

⁵Godsey and others (2005, 2011), Oviatt (2015), Miller (2016) for the timing of the occupation of the Provo shoreline and subsequent regression of Lake Bonneville to near Great Salt Lake level. Alternatively, data in Godsey and others (2005) may suggest that regression began earlier, shortly after 16.5 cal ka (see sample Beta-153158, with an age of 13,660 ± 50 1σ yr B.P. [16.5 cal ka] from 1.5 m below the Provo shoreline). Also, lacustrine carbonate deposits in caves reported by McGee and others (2012) seem to support an earlier Lake Bonneville regression beginning around 16.4 cal ka.

⁶Gilbert-episode highstand may have been very short lived; age represents lake culmination (Oviatt and others, 2005; Oviatt, 2014).

⁷Murchison (1989), Curry and James (1982)

⁸Miller and others (2005)

⁹Arnott and Stephens (1990)