

# CACHE VALLEY AQUIFER STORAGE AND RECOVERY— SITE ASSESSMENT FOR MILLVILLE CITY, CACHE COUNTY, UTAH

*by Paul Inkenbrandt*



**OPEN-FILE REPORT 636**  
**UTAH GEOLOGICAL SURVEY**  
*a division of*  
UTAH DEPARTMENT OF NATURAL RESOURCES  
**2014**

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*Cover photo: Millville, Utah*



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# CACHE VALLEY AQUIFER STORAGE AND RECOVERY—SITE ASSESSMENT FOR MILLVILLE CITY, CACHE COUNTY, UTAH

*by Paul Inkenbrandt*

## EXECUTIVE SUMMARY

Cache County is interested in pursuing aquifer storage and recovery (ASR) programs to store excess surface water in the Cache Valley principal aquifer. The City of Millville, located in a prime location for ASR, is having issues with elevated nitrate in the Glenridge well, a public water supply sourced from the Cache Valley principal aquifer. To initiate a small-scale ASR project and alleviate high nitrate, the city performed an initial injection and pumping test using the Glenridge well. Millville injected water from Garr Spring, another public water supply source of which they own water rights, into the Glenridge well for one week at a rate of 500 gallons per minute. Garr Spring water has an average nitrate concentration of 0.8 mg/l nitrate as nitrogen (Utah Division of Drinking Water, 2014). They then pumped the well while monitoring geochemistry to determine the effects on the Cache Valley principal aquifer system.

Results of the test are preliminary and show decreased nitrate values in the Glenridge well. While the increase in potentiometric surface was not precisely measured, it is likely small and widespread due to the high transmissivity of the aquifer, which was determined to be 135,000 ft<sup>2</sup>/day (12,540 m<sup>2</sup>/day). The pre-injection nitrate concentration in the Glenridge well was 7.65 mg/l nitrate as nitrogen, and the nitrate concentration after pumping more than 172% of the volume of water injected was 6.52 mg/l nitrate as nitrogen. There is likely some dispersion of the injected spring water via advection in the aquifer. Preliminary results indicate that the nitrate in the aquifer is stable and not reacting (it seems chemically conservative), but reaction rates have not been considered. A better understanding of prolonged injection is recommended before a full-scale ASR project is initiated.

## BACKGROUND

### Problem

Cache County leaders have expressed interest in storing water in aquifers of Cache Valley (figure 1). Based on this interest, the Utah Geological Survey (UGS) conducted an evaluation of potential aquifer storage and recovery (ASR) sites in Cache Valley (Thomas and others, 2011) and examined a gravel pit

site north of Logan, Utah (Inkenbrandt and others, 2013). In fall 2012, Millville City approached the UGS to express their interest and ability to participate in an ASR study involving the city's public water sources. They proposed injecting excess spring flow from Garr Spring into the Glenridge well. Millville possesses water rights for Garr Spring and owns Glenridge well, both of which are municipal water supply sources.

While Millville would like to store excess water in the Cache Valley principal aquifer, their greatest interest is diluting high nitrate concentrations within the aquifer. Nitrate values in the Glenridge well have risen by about 4 mg/l over the past 20 years, which correlates with an increase in the population of Millville (figure 2). Under the U.S. Environmental Protection Agency (U.S. EPA) regulations, nitrate cannot exceed a maximum contaminant level (MCL) of 10 mg/l (nitrate as nitrogen) (U.S. EPA, 2014). There are four major potential sources of nitrate for Millville to consider, none of which is exclusive of the others: (1) septic effluent from septic systems, (2) fertilizer, (3) livestock excrement, and (4) geologic sources (Lowe and Wallace, 2001; Roadcap and others, 2002).

### Objective

The objectives of this study are to: (1) determine if the Glenridge well and Garr Spring are suitable for conducting aquifer storage and recovery in the Cache Valley principal aquifer, and (2) determine the potential sources of nitrate contamination in the principal aquifer. As part of the objectives, I want to ensure that injecting water into the Glenridge well will not be detrimental to the aquifer.

### Millville City

Millville is located in southeast Cache Valley, in Cache County Utah, and has a population of 1869 (U.S. Census, 2014). Millville's Utah Division of Drinking Water system number is 03012, and a majority (>90%) of its water use is for domestic purposes (Utah Division of Water Rights, 2014). Millville operates four water sources, three of which—Garr Spring, the Glenridge well, and the Park well (figure 3)—are used as public drinking water sources (Utah Division of Water Rights, 2014).

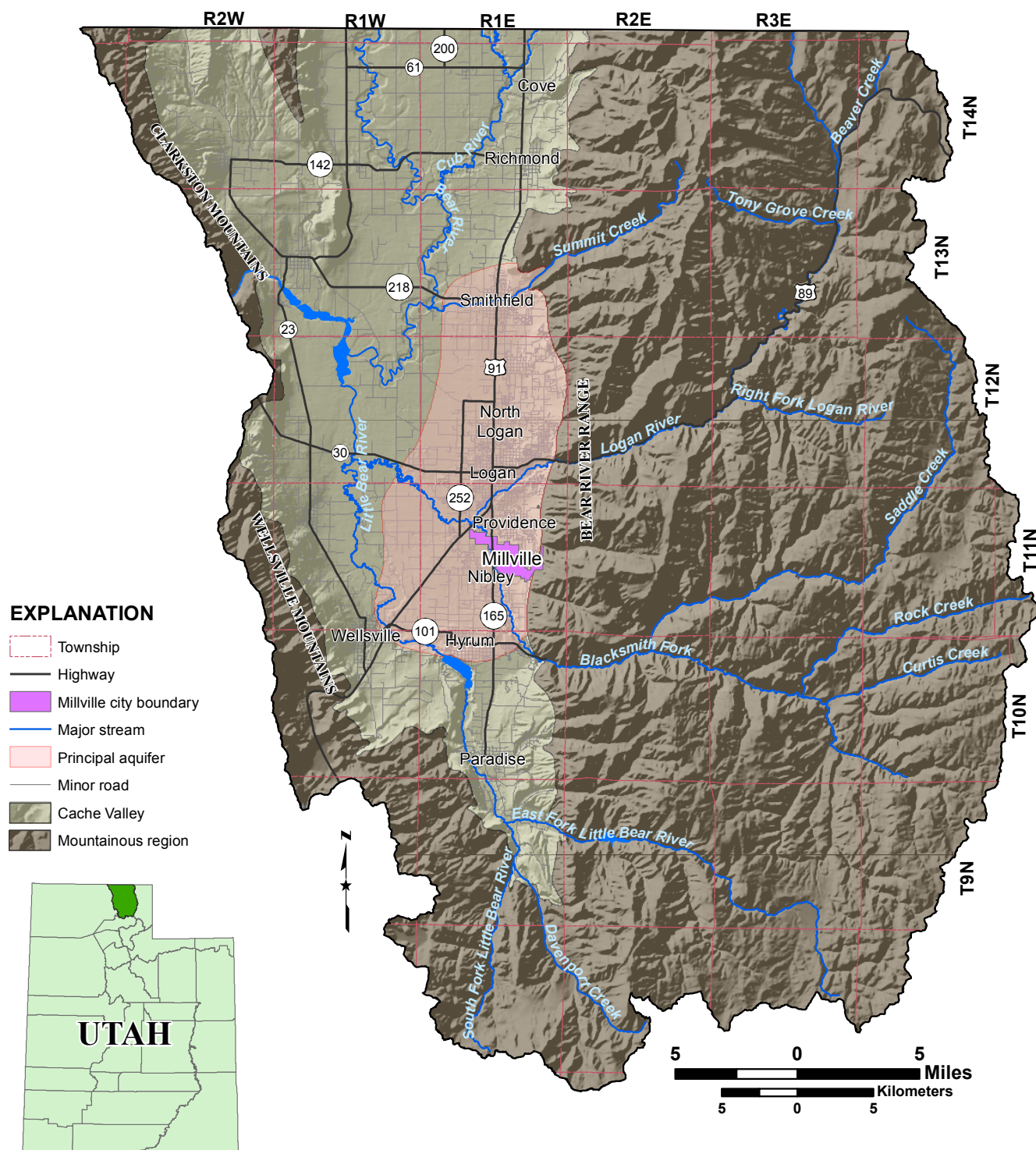
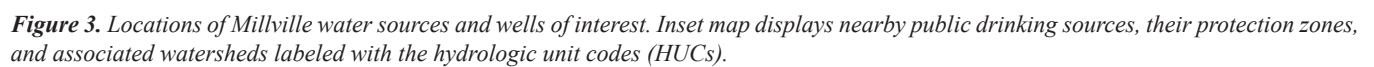


Figure 1. Location of the study area in Cache County, Utah.





Millville is currently an unsewered community, and each property has its own septic system. Like most communities in Cache Valley, Millville was historically an agricultural community (Millville History Book Committee, 1990), and animal operations that are still active include dairies, a mink farm, and livestock corrals.

### Garr Spring

The collection box of Garr Spring is in the foothills of the Bear River Range and is located at UTM Zone 12 coordinates 4613780 m North and 433079 m East (North American Datum 1983) at a surface elevation of 4849 feet (1478 m) above mean sea level. Millville operates a storage tank immediately west and downhill of the collection area. The area to the east, hydrologically upgradient of Garr Spring, is predominantly undeveloped U.S. Forest Service property.

In 2013, Millville used 524 acre-feet (ac-ft) of water, including irrigation water. About 36% (191 ac-ft) of that came from Garr Spring (Utah Division of Water Rights, 2014), which Mundorff (1971) reported as having relatively fresh water (specific conductance = 465 uS/cm; TDS = 258 mg/l) and a discharge of about 3.5 cubic feet per second (cfs) (2535 ac-ft/yr). Garr Spring water has an average nitrate concentration of 0.8 mg/l of nitrate as nitrogen (Utah Division of Drinking Water, 2014). Peterson (1946) reported the discharge of the spring as 5 cfs (3622 ac-ft/yr), whereas Beer (1967) measured a discharge of 4 cfs (2897 ac-ft/yr). Millville has 1.139 cfs (825 ac-ft/yr) in water rights from the spring (water right numbers 25-3510, 25-3069, 25-5170, 25-8394), which extend from October 1 to March 31 (181 days). The other water right holder on Garr Spring is Garr Spring Water Company (water right number 25-4528), which has 4.133 cfs (2994 ac-ft/yr) in water rights. Millville owns shares in the Garr Spring Water Company.

### Glenridge Well

The Glenridge well is near the center of Millville and is located at UTM Zone 12 4615423 m North and 431914 m East (North American Datum 1983) at a surface elevation of 4675 feet (1425 m) above mean sea level. The Utah Division of Water Rights well identification number (WIN) for the Glenridge well is 2722. Based on the well driller's report and construction information (appendix A), the Glenridge well is 385 feet deep with a 10-inch diameter steel casing and perforations from 269 to 369 feet. The depth to water from ground surface in 1972, the time of drilling, was 180 feet. In 2013, about 12% (61.6 ac-ft) of water used by Millville (524 ac-ft; includes irrigation) came from the Glenridge well (Utah Division of Water Rights, 2014a).

The water quality of the Glenridge well meets drinking water standards, but the nitrate-nitrogen levels are nearing the U.S. EPA MCL of 10.0 mg/l (nitrate as nitrogen) (figure 2).

The total dissolved solids concentration of the Glenridge well is 387 mg/l (Utah Division of Drinking Water, 2014). Water right number 25-5171 allots 2 cfs (1449 ac-ft/yr) of water to Millville from the Glenridge well (Utah Division of Water Rights, 2014).

### Park Well

The Park well is located near Garr Spring along the southeast border of Millville at UTM Zone 12 4613811 m North and 432552 m East (North American Datum 1983), and is at a surface elevation of 4705 feet (1434 m) above mean sea level. The Utah Division of Water Rights WIN for the Park well is 2721. In 2013, about 46% (240 ac-ft) of water used by Millville came from the Park well (Utah Division of Water Rights, 2014). Water from this well is relatively fresh, with an average nitrate concentration of 0.73 mg/l nitrate as nitrogen and an average total dissolved solids concentration of 267 mg/l (Utah Division of Drinking Water, 2014). The Park well was drilled in 1976 to a depth of 398 feet and has a 12-inch diameter steel casing. See appendix A for more details of the Park well.

### Aquifer Storage and Recovery

Aquifer storage and recovery (ASR), or conjunctive use, is the method of storing water in an aquifer when the water is available and recovering that water when needed (Pyne, 2005). While groundwater recharge generally occurs naturally, the recharge aspect of ASR is usually induced via human intervention and is commonly referred to as "managed aquifer recharge" (MAR) (Pyne, 2005). MAR is conducted either using a surface recharge basin, as examined by Inkenbrandt and others (2013), or by injection into a well (Pyne, 2005), which is the case for this study.

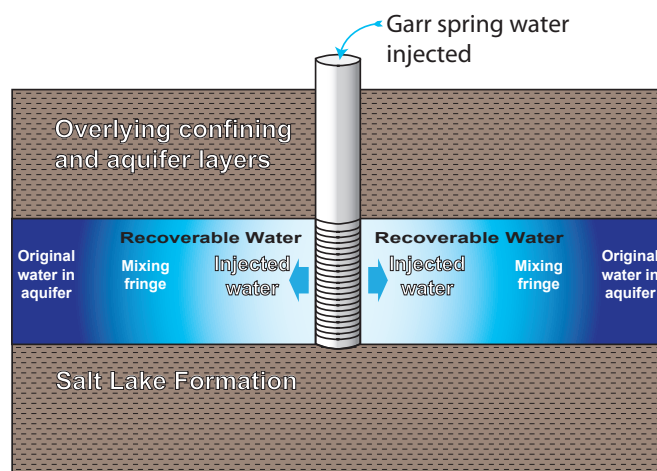
An injection well offers the advantage of injecting water directly into the target aquifer without encountering confining layers or other impedances that could exist between a surface recharge basin and the target aquifer. In ASR configurations, wells can be used solely for injection or for both injection and pumping. The advantages of using a "dual purpose" well for both injection and extraction are that only one well is required for the operation and the pump in the well can be used to help remediate clogging of the well screen—a major issue with ASR injection wells (Pyne, 2005).

Injecting water into an aquifer system creates an effective "bubble" of the injected water in the aquifer (figure 4). Natural groundwater flow and diffusion disperse that bubble over time, but some quantity of the injected water can usually be retrieved. The proportion of injected water that can be retrieved is known as recovery efficiency.

While stored in the aquifer, the injectate, native aquifer water, and the solid aquifer material can undergo hydrogeochemi-

cal changes depending on the relative difference between the chemistries of the two waters. If the aquifer is dominated by reducing conditions and an oxygen-rich injectate is introduced, the chemistry of the injectate water could change the oxidation state of ions in the native aquifer water and/or the aquifer material, resulting in the potential mobilization of ions (figure 5a). Arsenic mobilization is a common issue (figure 5b), via oxidation of arsenopyrite in the aquifer. Other concerns include mobilization of uranium, mercury, nickel, chromium, cobalt, and zinc. Microbiota often play an important role in the mobilization and demobilization of these and other ions, especially in the case of nitrate. A primary concern for the water surrounding the Glenridge well is potentially mobilizing nitrate via the oxidation of nitrite and ammonia (nitrification) (figure 5c). While the U.S. EPA does currently not regulate sulfate, the presence or absence of sulfate can also be indicative of reducing conditions (figure 5d).

Nitrate can also potentially be mobilized by raising the water table of an unconfined aquifer, as observed by Nishikawa and others (2003). If the aquifer is unconfined and the nitrate source is near the water table, then raising the groundwater level would result in greater contamination of the aquifer water. This is especially true if the recharge source is a recharge basin, as the water will travel through the contaminated unsaturated zone before reaching the saturated zone (figure 6). The geologic setting examined by Nishikawa and others (2003) is very similar to that of Millville, with the important exception of faults dissecting the Cache-Valley basin-fill aquifer into different aquifers.



**Figure 4.** Conceptual diagram of injected-water storage in the Cache Valley principal aquifer.

## Regulation

Because water is being injected into a drinking-water aquifer, the injection well used for this study (Millville's Glenridge well) is subject to regulation by the Underground Injection Control (UIC) Program of the Utah Division of Water Quality and the Utah Division of Water Rights (Utah State Legislature,

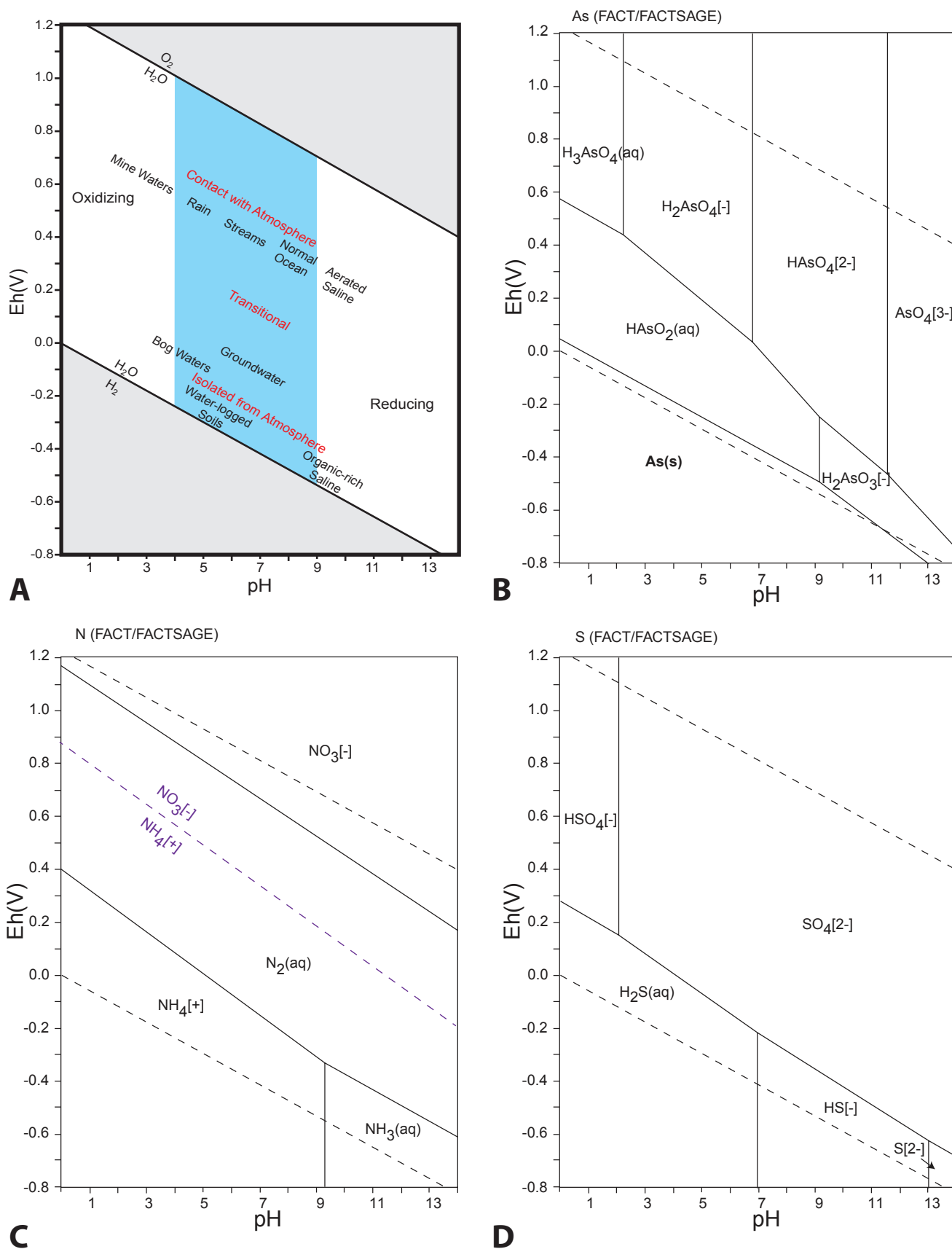
2014). As a public supply well, it is also subject to the regulatory criteria of the Utah Division of Drinking Water. Any entity injecting water in the state of Utah must follow rules outlined by Title 73, Chapter 3b of the Utah State Code (Utah State Legislature, 2014). The UIC Program classifies ASR wells as Class 5B4 injections wells, which are used to replenish water in an aquifer for subsequent use (Utah Division of Water Quality, 2014). For an entity to inject water into a Class 5B4 well, a permit application must be filed with the UIC Program (appendix B). The entity must also file an application with the Utah Division of Water Rights to both recharge and recover water (appendix C). State entities overseeing the aquifer storage and recovery process either require or have great interest in a hydrologic study defining: (1) the area of the aquifer impacted by injection, (2) implications of injection of foreign water into the groundwater system, (3) the hydrogeology of the area, and (4) the capabilities of the entity injecting water.

## Hydrogeology

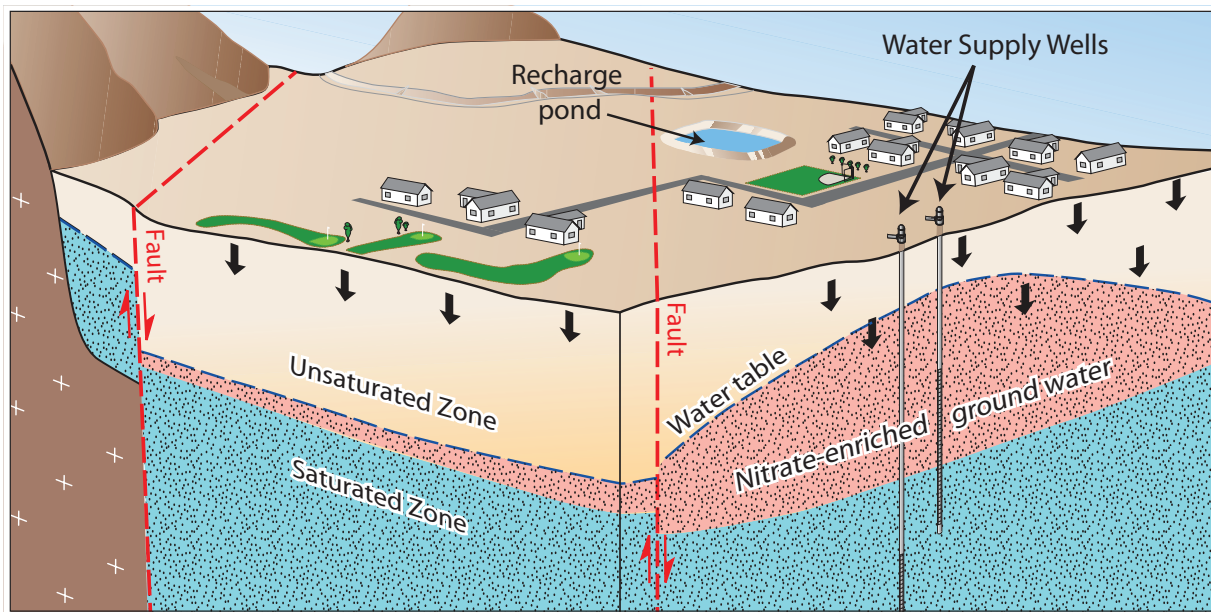
The Cache Valley principal aquifer system (figure 1), the primary aquifer for drinking-water supplies, is a complex multiple-aquifer system composed of basin fill under both unconfined and confined conditions (Bjorklund and McGreevy, 1971; Kariya and others, 1994). The basin fill is unconsolidated sediment consisting of silt, sand, and gravel, which were deposited in fluvial, alluvial fan, landslide, and near-shore lacustrine environments. Each layer is bounded by layers of silt and clay primarily deposited by offshore lacustrine environments (Bjorklund and McGreevy, 1971; Lowe, 1987). The basin fill is more than several hundred feet thick at many locations along the valley center (Kariya and others, 1994).

Bjorklund and McGreevy (1971) concluded that groundwater in the principal aquifer is unconfined along the margins of Cache Valley, but is confined in many areas toward the center of the valley where many flowing wells exist. Using over 200 well drillers' logs, isotopic signatures, and carbon-14 age estimates, Robinson (1999) developed and Olsen (2007) improved a conceptual model (figure 7) of the Cache Valley principal aquifer system. Robinson (1999) closely examined the aquifer system in southern Cache Valley to the west and south of Millville and reported that groundwater in the aquifer is relatively old and slow moving.

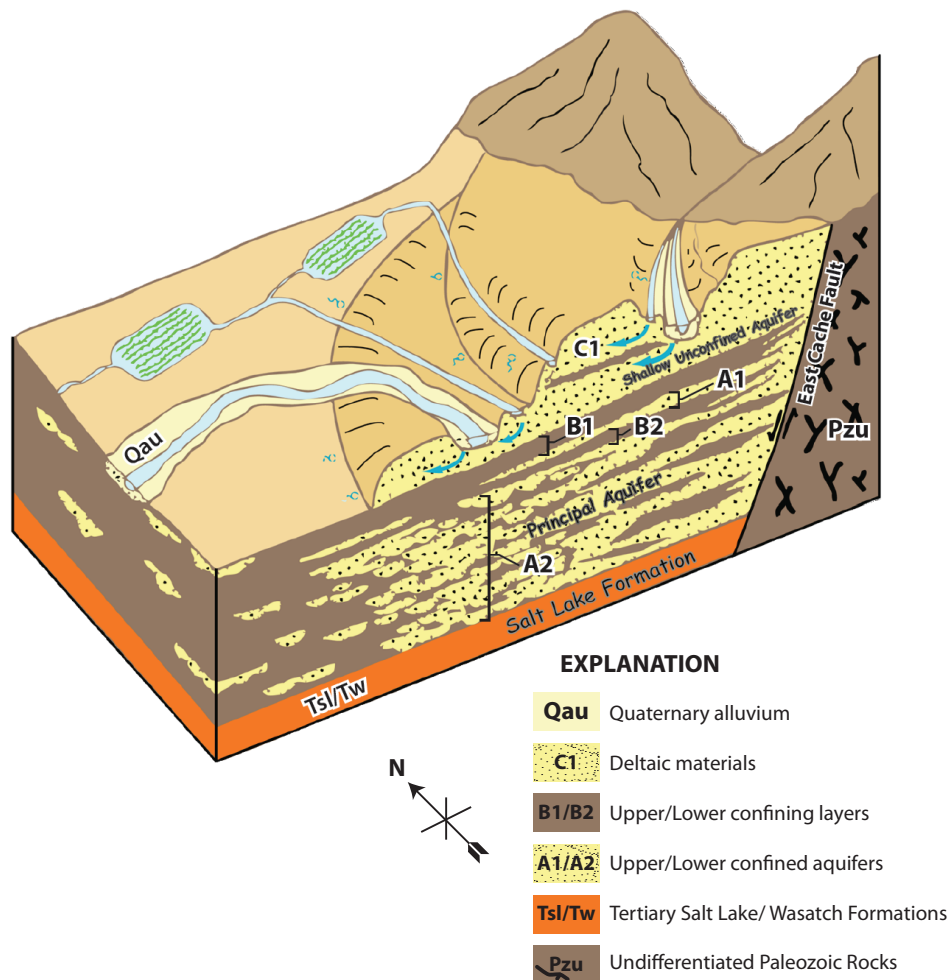
The boundary between unconfined and confined conditions is gradational near the margins of the basin, which is within a couple miles of the basin-bounding fault and where leaky conditions may exist. The presence or absence of continuous clay layers and the vertical hydrologic gradient dictate where groundwater recharges and discharges. Based on mapping conducted by Anderson and others (1994), groundwater recharge occurs mainly at the margins of Cache Valley, while discharge is predominantly near the center of the valley (figure 8). Using well drillers' logs and the hydrologic gradient, Anderson and others (1994) subdivided the Cache Valley



**Figure 5.** Pourbaix (Eh-pH) diagrams of (A) natural waters (modified from Garrels and Christ, 1965), (B) arsenic (from Takeno, 2005), (C) nitrogen (modified from Takeno, 2005), and (D) sulfur (from Takeno, 2005).

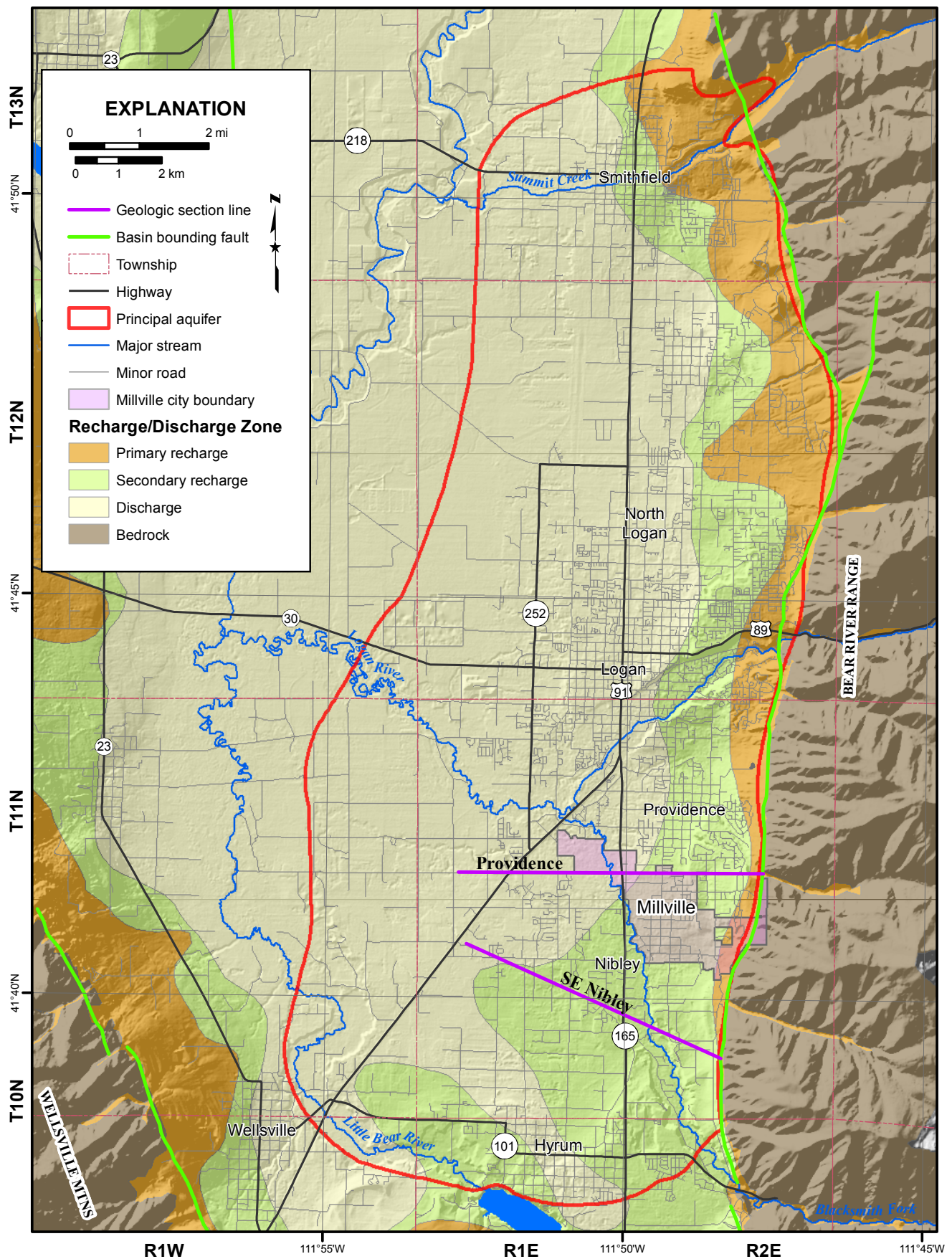


**Figure 6.** Conceptual model of water from the recharge pond infiltrating through an area with a high density of septic systems, creating a recharge mound and mobilizing nitrate in an area previously unsaturated (modified from Nishikawa and others, 2003).



**Figure 7.** Conceptual block diagram of Cache Valley hydrostratigraphic units (modified from Olsen, 2007) as defined by Robinson (1999). This conceptual block diagram best represents the area near Logan, Utah.





**Figure 8.** Recharge and discharge areas in Cache Valley (Anderson and others, 1994). See figure 9 for the geologic cross sections (purple lines).

basin-fill into three categories: (1) primary recharge—less than 20 feet of clay and a downward hydrologic gradient, (2) secondary recharge—confining layers (greater than 20 feet) and a downward hydrologic gradient, and (3) discharge areas—upward hydrologic gradient.

When selecting potential candidate areas for ASR surface-spreading sites (areas where surface water can infiltrate into the ground), Thomas and others (2011) created several geologic cross sections to gain a better understanding of the continuity of clay layers and aquifer systems within the principal aquifer. Two of those sections bracket the area near Millville (figure 9).

## Groundwater Chemistry

Lowe and others (2003) sampled 165 wells for a study on groundwater quality in Cache Valley and created recommendations for septic tank densities based on their findings. Lowe and others (2003) recommended a maximum of one-third septic system per acre for the Millville area. Using data collected in 1997 by Lowe and others (2003), Robinson (1999) evaluated the southern Cache Valley aquifer system. Robinson (1999) interpolated sulfate values for the southern portion of the valley and found high levels of sulfate near Blacksmith Fork (figure 10).

As previously stated, nitrate contamination near the Millville Glenridge well has increased over time approaching the U.S. EPA MCL of 10 mg/l (nitrate as nitrogen). Nitrate concentration in the Glenridge well is high compared to concentrations at Garr Spring and the Park well, which are at or near 1 mg/l (Utah Division of Drinking Water, 2014). As of the publication of this report, one well in Cache Valley (Mendon, Utah) has been shut down due to nitrate contamination.

## METHODS

### Approach

To determine the nitrate source and the suitability of ASR for Millville, we: (1) collected baseline data for the aquifer system, (2) conducted an injection test on the Glenridge well using Garr Spring water, (3) conducted a pumping test on the Glenridge well, and (4) modeled potential impacts on the aquifer system.

I compiled existing geochemical data and sampled groundwater for stable nitrogen, oxygen, and hydrogen isotopes, chemicals from septic systems, and standard major isotopes (including nitrate). In our groundwater-level measurement approach, I compiled water levels, conducted a microgravity survey, and Millville City conducted an injection test and pumping test using the Glenridge well. Based on the infor-

mation I gained from the geochemistry and water-level data, I revised the drinking water source protection zone for the Glenridge well and predicted future behavior of the aquifer system if subjected to proposed injections. With the revised drinking water source protection zone, I identified additional potential contaminant sources near Millville.

## Background Measurements

### Potential Nitrate Sources

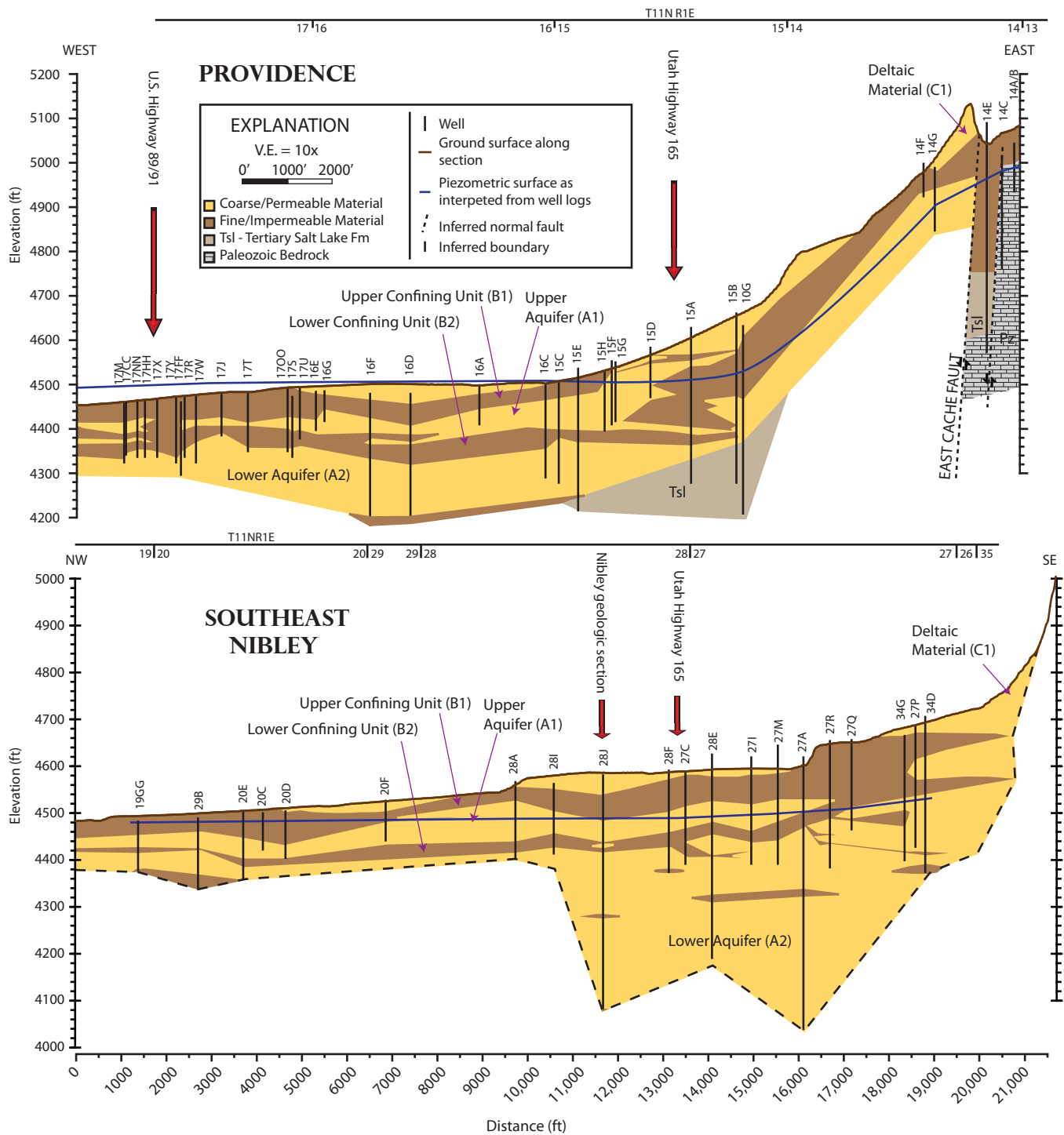
A potential contaminant map showing the location of septic systems and animal operations was created using the Cache County parcel map of Millville (Utah Automated Geographic Reference Center, 2013a). Points were created using the centroid of each parcel. A few older parcels in Providence suspected of having septic systems, as reported in informal communications with Providence city workers, were included. I then generated a point density map from the centroid points. To include potential animal-related sources of nitrate, I used air photos and water-related land use maps (Utah Automated Geographic Reference Center, 2013b) to determine land use change over time.

### Groundwater Levels

To determine groundwater flow direction and gradient, which are important components of Drinking Water Source Protection (DWSP) delineation, contaminant transport, and injection modeling, I created a high-resolution potentiometric surface map using ArcGIS software. A combination of compiled and field-collected data (appendix D) were used to generate the map. I used 30-foot (10m) horizontal resolution digital elevation models (DEM) from the National Elevation Dataset (NED) (USGS, 2014a) to assign land-surface elevations for wells that I did not measure in the field.

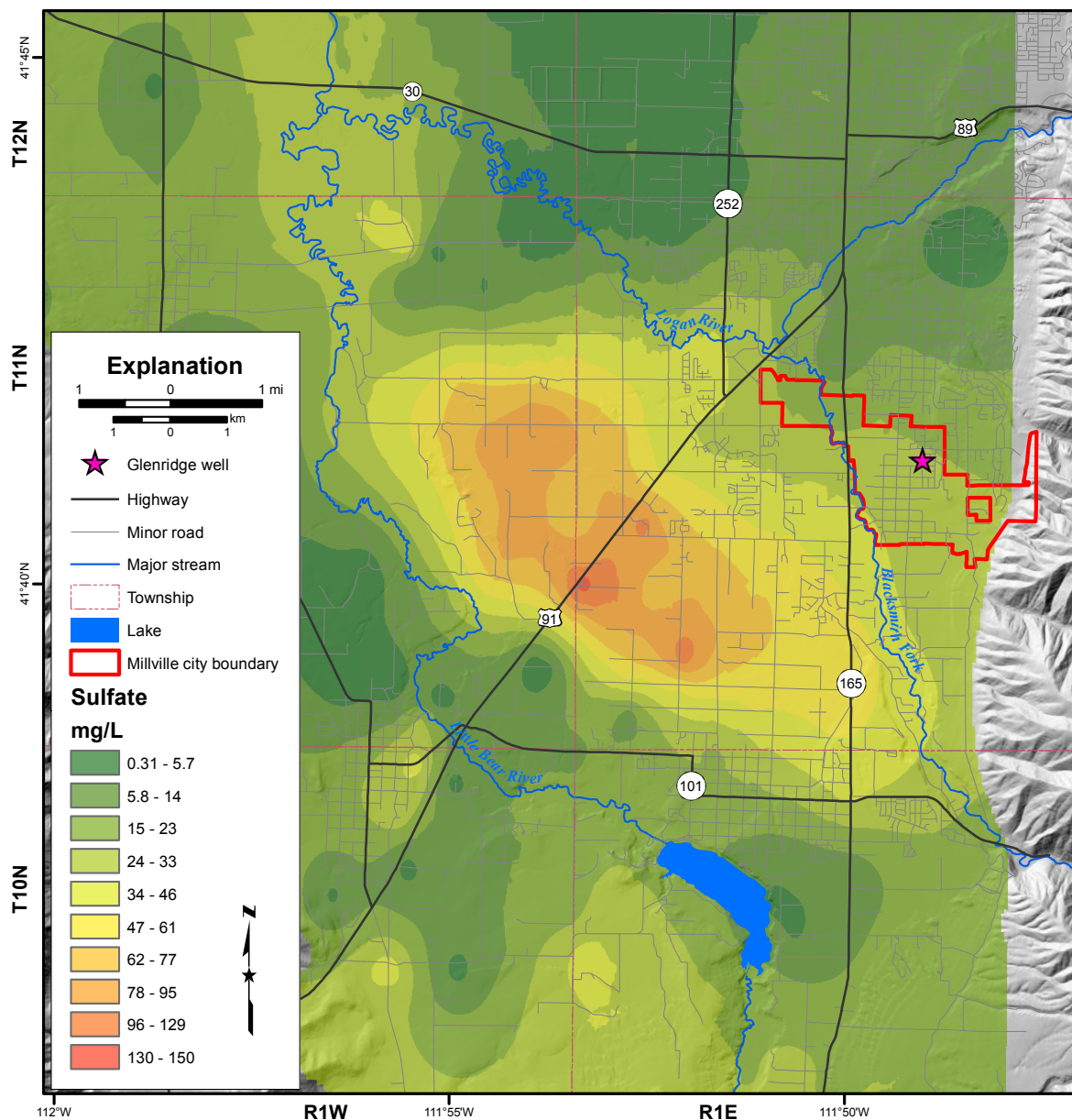
To collect field data, I used a high-resolution Trimble real-time kinetic global positioning system in combination with the Utah Reference Network (TURN) (Utah Automated Geographic Reference Center, 2013c) to measure well-casing-top elevations and a Solinst water level sounder to measure depth to water.

I supplemented field measurements with compiled data from the U.S. Geological Survey (USGS) and the Utah Division of Water Rights, and I compiled existing data from the USGS NWIS database (USGS, 2014b). For wells having multiple measurements, an average of depth-to-water data was calculated. I determined that the averaging of measurements was appropriate as I observed no significant long-term trend in water levels for Cache Valley, and I wanted to represent the long-term static groundwater level. I also compiled data from the Utah Division of Water Rights points of diversion (WR-POD) shapefile and associated tables.



**Figure 9.** Geologic cross sections of the Millville area (Thomas and others, 2011). See figure 8 for section locations. Elevation is in feet above mean sea level.





**Figure 10.** Sulfate concentrations in the Millville area (modified from Robinson, 1999). Note the high concentration of sulfate downgradient (west) of Blacksmith Fork.

The cokriging gridding method was used to interpolate the groundwater level elevation data. I used NED data (USGS, 2014a) and 645 groundwater-level measurements compiled from 576 wells in the principal aquifer area for the interpolation. I used the entire principal aquifer area to better model spatial relationships between groundwater levels. Cokriging is a geostatistical method that assumes one can improve estimates of a value of a variable in space if it is spatially dependent on other variables (ESRI, 2014). The advantage of using a geostatistical method is it allows for error estimates in interpolation (ESRI, 2014).

I determined gradient and flow direction by generating slope and aspect rasters of the groundwater level interpolation. I averaged the aspect and slope orientations within a two-mile radius of the well and west of the East Cache Valley bounding fault.

### Hydrogeology

Using data from 60 well drillers' logs (appendix E), I created two hydrogeologic cross sections of the Millville area. I simplified drillers' unit descriptions to "high" and "low" permeability. When permeability information wasn't available, I labeled clay-bearing units as "low" permeability and



sand, gravel, and conglomerate units as “high” permeability. To map hydrogeologic units at the surface, I interpreted permeability from the Quaternary units presented in the Logan 1:24,000-scale geologic map (Evans and others 1996).

The same wells were used to determine total clay thickness within the first 150 feet of the subsurface. I chose 150 feet as the cutoff because this is above the depth to water in most of the Millville area. For each well, I summed the thicknesses of layers labeled “clay” by the drillers in the first 150 feet below ground surface. In order to determine the coverage of clay in the area, I used the Empirical Bayesian Kriging (EBK) method to interpolate total clay thickness between wells. EBK is a geostatistical interpolation method used to automatically generate localized geostatistical models that approximate spatial variation in a variable (ESRI, 2014).

## Hydrogeochemistry

Four wells and two springs were sampled (figure 3) for nitrate, and a subset of those sites were sampled for stable oxygen and hydrogen isotopes in water, arsenic, metals, and general chemistry. The Utah State University Water Laboratory conducted a preliminary analysis on a water sample from the Glenridge well for caffeine and a suite of other chemicals commonly sourced from septic system. Presence of these chemicals could indicate septic contribution. Prior to conducting the injection test, I was required by the Department of Environmental Quality to sample for a wide variety of constituents (appendix F).

Prior to the project, Gary Larsen (Millville City public works director) collected nitrate data from several sites over several years. The Utah State Health Department Laboratory analyzed all samples, except for stable isotopes, which were analyzed at the Utah State University Geology Department Isotope Lab.

Stable isotopes of nitrogen and oxygen in nitrate can be used to narrow down potential sources of nitrate in groundwater. Different sources plot in different regions on a graph of the ratios of oxygen and nitrogen isotopes, with some overlap (Clark and Fritz, 1997). The locations of the samples on the graph give some indication as to the source of nitrate. If the samples fall into the overlap areas, the source is mixed or ambiguous (Clark and Fritz, 1997). Waterloo Laboratory analyzed a water sample from the Glenridge well for nitrogen and oxygen isotope ratios in nitrate to determine a possible nitrate source.

I also used stable isotopes of oxygen and hydrogen in water, as well as nitrate concentrations, to determine the “make-up” of the post-injection water extracted from the Glenridge well. Based on ratios of source concentration to point-measurement concentrations, assuming nitrate and stable isotopes are conservative, I determined the percentage of pumped water that is original to the Cache Valley principal aquifer (i.e., non-in-

jected water) relative to the amount of water pumped. Using EBK interpolation on the most recent nitrate values, I created a nitrate concentration map to determine extent and orientation of the nitrate plume.

## Injection and Pumping Test

### Injection Configuration

Water from Garr Spring was injected through existing city water lines into the pump shaft of the Glenridge well. Before injection into the well, the pump was removed and check valves and flow meters were reversed to allow water to flow opposite of the typical (pumping) flow direction. A 4-inch diameter injection pipe with a 1.625-inch diameter orifice at the end was set to 210 feet below ground surface, which is 30 feet below static water level.

Injection began at 10:45 a.m. on 3/10/2014 and continued until 3/17/2014 at 10:51 a.m. (table 1) at a mean rate of 476 gallons per minute (gpm) (1800 liters per minute). I measured injection rate and volume using an in-pipe volume meter. The meter requires the pipe to be full to be accurate, which was the case during the injection test. Millville injected a total of 4,987,000 gallons (15.3 ac-ft; 19,000 m<sup>3</sup>) of Garr Spring water into the Glenridge well over a period of 7 days (table 1).

### Pumping Configuration

After injecting was complete, the pump was placed back into the well with the pump intake at 231 feet below ground surface. Pumping began at 9:15 am on 3/19/2015 and continued until 3/24/2014 at 10:56 a.m. at a mean rate of 290 gpm. I measured the pumping rate and volume using the same flow-meter used for injection. Millville pumped a total of 2,117,000 gallons (6.5 ac-ft; 8000 m<sup>3</sup>) from Glenridge well over a period of 5 days (table 2). The well sat idle until 5/5/2014 when the pump was turned on to purge and sample the well. The well pump was turned on for extended municipal use on 7/1/2014.

### Measurements

During the test, I measured water levels with a combination of water level sounders and pressure transducers. Using the pressure transducers, I measured at a minimum frequency of 1 sample every 10 minutes and a maximum frequency of 1 sample every 10 seconds. I increased measurement frequency during the beginning and the end of injection and pumping to capture rapid changes in well water level. Due to well access issues, I was not able to continuously record water levels in the Glenridge well prior to pumping to analyze for antecedent trends. However, I periodically manually measured groundwater levels prior to pumping and extended the duration of post-pumping measurement to examine the well for other water-level-changing influences. See appendix G for a table of well-water level measurements.

**Table 1.** Injection record of Glenridge well.

Date-Time	Pipe Meter Reading (gallons)	Total Water Injected (gallons)	Total Water Injected (acre-ft)	Injection Rate (gpm)	Time Since Injection Started (days)
3/10/2014 10:45	393235000				
3/10/2014 11:02	393239250	4250	0.013	250	0.0118
3/10/2014 11:04	393241600	6600	0.0203	810.34	0.0138
3/10/2014 11:33	393254725	19725	0.0605	463.78	0.0335
3/10/2014 12:16	393276300	41300	0.1267	500	0.0634
3/10/2014 14:42	393349000	114000	0.3499	498.86	0.1646
3/10/2014 14:46	393351300	116300	0.3569	498.19	0.1678
3/10/2014 14:56	393356200	121200	0.3719	500	0.1747
3/10/2014 17:16	393426000	191000	0.5862	498.69	0.2719
3/10/2014 17:17	393426500	191500	0.5877	500	0.2725
3/10/2014 17:19	393427500	192500	0.5908	495.87	0.2739
3/11/2014 7:57	393864000	629000	1.9303	497.17	0.8836
3/11/2014 7:59	393865000	630000	1.9334	500	0.885
3/11/2014 11:43	393976000	741000	2.274	496.09	1.0404
3/11/2014 11:44	393976500	741500	2.2756	491.8	1.0411
3/11/2014 15:44	394095500	860500	2.6408	495.97	1.2077
3/11/2014 15:45	394096000	861000	2.6423	491.8	1.2084
3/11/2014 15:46	394096500	861500	2.6438	500	1.2091
3/11/2014 17:57	394161500	926500	2.8433	495.49	1.3002
3/11/2014 17:58	394162000	927000	2.8449	491.8	1.3009
3/11/2014 18:08	394167000	932000	2.8602	497.51	1.3079
3/12/2014 7:53	394575500	1340500	4.1138	494.93	1.8811
3/12/2014 7:54	394576000	1341000	4.1154	491.8	1.8818
3/13/2014 7:10	395266000	2031000	6.2329	494.55	2.8507
3/13/2014 15:15	395505000	2270000	6.9664	492.78	3.1875
3/13/2014 19:00	395613000	2378000	7.2978	480	3.3438
3/14/2014 7:15	395980000	2745000	8.4241	499.32	3.8542
3/14/2014 15:30	396229000	2994000	9.1882	503.03	4.1979
3/14/2014 20:00	396362000	3127000	9.5964	492.59	4.3854
3/15/2014 10:30	396793000	3558000	10.9191	495.4	4.9896
3/16/2014 10:00	397477000	4242000	13.0182	485.11	5.9688
3/16/2014 18:30	397738000	4503000	13.8192	511.76	6.3229
3/17/2014 9:46	398190500	4955500	15.2079	493.77	6.9593
3/17/2014 9:47	398191000	4956000	15.2094	500	6.96
3/17/2014 10:43	398218500	4983500	15.2938	494.6	6.9986
3/17/2014 10:46	398220100	4985100	15.2987	477.61	7.0009
3/17/2014 10:46	398220200	4985200	15.299	428.57	7.0011
3/17/2014 10:46	398220300	4985300	15.2993	428.57	7.0013
3/17/2014 10:47	398220400	4985400	15.2996	500	7.0014
3/17/2014 10:47	398220500	4985500	15.2999	461.54	7.0016
3/17/2014 10:47	398220600	4985600	15.3002	461.54	7.0017
3/17/2014 10:47	398220700	4985700	15.3005	461.54	7.0019
3/17/2014 10:47	398220800	4985800	15.3008	461.54	7.002
3/17/2014 10:48	398220900	4985900	15.3011	500	7.0022
3/17/2014 10:48	398221000	4986000	15.3015	428.57	7.0023
3/17/2014 10:49	398221300	4986300	15.3024	418.6	7.0028
3/17/2014 10:49	398221400	4986400	15.3027	500	7.003
3/17/2014 10:49	398221500	4986500	15.303	461.54	7.0031
3/17/2014 10:49	398221600	4986600	15.3033	428.57	7.0033
3/17/2014 10:49	398221700	4986700	15.3036	461.54	7.0034
3/17/2014 10:50	398221800	4986800	15.3039	400	7.0036
3/17/2014 10:50	398221900	4986900	15.3042	428.57	7.0038
3/17/2014 10:50	398222100	4987100	15.3048	428.57	7.0041
3/17/2014 10:51	398222200	4987200	15.3051	375	7.0043
3/17/2014 10:51	398222300	4987300	15.3054	352.94	7.0045
3/17/2014 10:51	398222400	4987400	15.3057	428.57	7.0046

Table 2. Pumping record of Glenridge well.

Date-Time	Pipe Meter Reading (gallons)	Total Volume Pumped (acre-feet)	Time Since Pumping Started (days)	Total Volume Pumped (gallons)	Pumping Flow Rate (gpm)	Volume Pumped Relative to Volume Injected
3/19/2014 9:15	398221530					
3/19/2014 10:18	398221530		0.0000		0.0	
3/19/2014 10:24	398223500	0.006	0.0046	1970	295.5	0.04%
3/19/2014 10:25	398223600	0.0064	0.0049	2070	300.0	0.04%
3/19/2014 10:25	398223700	0.0067	0.0051	2170	315.8	0.04%
3/19/2014 10:25	398223800	0.007	0.0053	2270	300.0	0.05%
3/19/2014 10:26	398224000	0.0076	0.0058	2470	279.1	0.05%
3/19/2014 10:26	398224100	0.0079	0.0061	2570	285.7	0.05%
3/19/2014 10:27	398224200	0.0082	0.0063	2670	300.0	0.05%
3/19/2014 10:27	398224300	0.0085	0.0065	2770	272.7	0.06%
3/19/2014 10:27	398224400	0.0088	0.0068	2870	285.7	0.06%
3/19/2014 10:28	398224700	0.0097	0.0075	3170	305.1	0.06%
3/19/2014 10:29	398224800	0.01	0.0077	3270	285.7	0.07%
3/19/2014 10:31	398225350	0.0117	0.0090	3820	289.5	0.08%
3/19/2014 10:31	398225400	0.0119	0.0092	3870	272.7	0.08%
3/19/2014 10:31	398225500	0.0122	0.0094	3970	315.8	0.08%
3/19/2014 11:18	398239300	0.0545	0.0422	17770	291.8	0.36%
3/19/2014 11:19	398239400	0.0548	0.0425	17870	285.7	0.36%
3/19/2014 13:20	398274900	0.1638	0.1268	53370	292.4	1.07%
3/19/2014 13:20	398275000	0.1641	0.1270	53470	285.7	1.07%
3/19/2014 13:51	398283800	0.1911	0.1479	62270	292.4	1.25%
3/19/2014 17:43	398350650	0.3963	0.3091	129120	288.1	2.59%
3/19/2014 22:22	398433400	0.6502	0.5031	211870	296.1	4.25%
3/19/2014 22:22	398433500	0.6505	0.5034	211970	285.7	4.25%
3/19/2014 22:23	398433600	0.6508	0.5036	212070	300.0	4.25%
3/19/2014 22:23	398433800	0.6514	0.5041	212270	292.7	4.26%
3/19/2014 22:33	398436500	0.6597	0.5105	214970	291.4	4.31%
3/20/2014 9:16	398625700	1.2404	0.9570	404170	294.3	8.10%
3/20/2014 9:39	398631500	1.2582	0.9731	409970	249.3	8.22%
3/20/2014 10:44	398650500	1.3165	1.0183	428970	292.5	8.60%
3/20/2014 10:44	398650700	1.3171	1.0187	429170	292.7	8.61%
3/20/2014 10:51	398652600	1.3229	1.0232	431070	292.3	8.64%
3/20/2014 10:51	398652700	1.3232	1.0235	431170	300.0	8.65%
3/20/2014 10:52	398652800	1.3235	1.0237	431270	272.7	8.65%
3/20/2014 10:52	398652900	1.3238	1.0239	431370	315.8	8.65%
3/21/2014 8:49	399038300	2.5066	1.9383	816770	292.7	16.38%
3/21/2014 8:49	399038500	2.5072	1.9388	816970	307.7	16.38%
3/21/2014 8:50	399038700	2.5078	1.9393	817170	292.7	16.38%
3/21/2014 8:51	399039000	2.5087	1.9400	817470	281.3	16.39%
3/21/2014 8:51	399039100	2.509	1.9402	817570	300.0	16.39%
3/21/2014 9:23	399048200	2.537	1.9618	826670	292.6	16.58%
3/21/2014 12:43	399106800	2.7168	2.1010	885270	292.4	17.75%
3/21/2014 14:28	399137500	2.811	2.1739	915970	292.5	18.37%
3/21/2014 14:36	399139800	2.8181	2.1793	918270	293.0	18.41%
3/21/2014 16:16	399169000	2.9077	2.2486	947470	292.6	19.00%
3/21/2014 16:16	399169100	2.908	2.2489	947570	285.7	19.00%
3/21/2014 16:16	399169200	2.9083	2.2491	947670	285.7	19.00%
3/21/2014 16:17	399169300	2.9086	2.2494	947770	285.7	19.00%
3/22/2014 7:45	399441200	3.743	2.8944	1219670	292.7	24.46%
3/22/2014 7:46	399441400	3.7436	2.8949	1219870	285.7	24.46%
3/22/2014 7:46	399441500	3.7439	2.8951	1219970	300.0	24.46%
3/24/2014 10:31	400332600	6.4786	5.0091	2111070	292.7	42.33%
3/24/2014 10:31	400332700	6.4789	5.0093	2111170	315.8	42.33%

Table 2. Continued

Date-Time	Pipe Meter Reading (gallons)	Total Volume Pumped (acre-feet)	Time Since Pumping Started (days)	Total Volume Pumped (gallons)	Pumping Flow Rate (gpm)	Volume Pumped Relative to Volume Injected
3/24/2014 10:48	400337800	6.4946	5.0214	2116270	292.5	42.43%
3/24/2014 10:49	400337900	6.4949	5.0217	2116370	260.9	42.43%
3/24/2014 10:49	400338000	6.4952	5.0219	2116470	260.9	42.44%
3/24/2014 10:50	400338100	6.4955	5.0222	2116570	240.0	42.44%
3/24/2014 10:50	400338150	6.4957	5.0223	2116620	272.7	42.44%
3/24/2014 10:50	400338200	6.4958	5.0225	2116670	187.5	42.44%
3/24/2014 10:50	400338250	6.496	5.0227	2116720	214.3	42.44%
3/24/2014 10:50	400338300	6.4961	5.0229	2116770	214.3	42.44%
3/24/2014 10:51	400338350	6.4963	5.0230	2116820	214.3	42.44%
3/24/2014 10:51	400338400	6.4964	5.0232	2116870	230.8	42.44%
3/24/2014 10:51	400338450	6.4966	5.0233	2116920	230.8	42.45%
3/24/2014 10:51	400338500	6.4967	5.0235	2116970	214.3	42.45%
3/24/2014 10:52	400338550	6.4969	5.0236	2117020	250.0	42.45%
3/24/2014 10:52	400338600	6.497	5.0238	2117070	214.3	42.45%
3/24/2014 10:52	400338650	6.4972	5.0240	2117120	200.0	42.45%
3/24/2014 10:52	400338700	6.4973	5.0241	2117170	187.5	42.45%
3/24/2014 10:53	400338750	6.4975	5.0243	2117220	187.5	42.45%
3/24/2014 10:53	400338800	6.4977	5.0245	2117270	187.5	42.45%
3/24/2014 10:56	400338950	6.4981	5.0264	2117420	55.6	42.46%
3/24/2014 10:56	400338950	6.4981	5.0264	2117420	0.0	42.46%

## Modeling

### Aquifer Test Analyses

I used AQTESOLV (Duffield, 2007) computer software to determine aquifer transmissivity from test data collected during the pumping period. I did not analyze the injection data. I applied a Theis (1935)/Hantush (1961) solution for a confined aquifer. Variation in the groundwater level data prevented a more precise analysis and model of the drawdown and recovery curves. I applied alternate solutions that produced results within an order of magnitude of the Theis (1935)/Hantush (1961) approximation.

### Injection Model

I used AQTESOLV (Duffield, 2007) to determine potentiometric surface changes due to injection and to create an area of influence (AOI). Using the parameters determined from aquifer test analyses, the groundwater level map, and geologic cross sections, I modeled the magnitude of potentiometric changes in the aquifer that various rates of injection would induce. For injection, I assumed that Millville would inject water into the aquifer system at 300 gallons per minute for 181 days per year and then pump out about 60 ac-ft/yr of that water during July, August, and September.

## Drinking Water Source Protection Zones

I reevaluated the DWSP zones based on recalculated aquifer properties and the presence of the East Cache fault. I determined horizontal hydrologic gradient using the groundwater level map created for this study. I used cross sections created for this study and the Glenridge well log (appendix A) to determine aquifer thickness.

To determine hydraulic conductivity, I recalculated data from Inkenbrandt (2010) using information gained from the aquifer test analyses and incorporated more recent data collected during the pumping test. I used the EBK method to interpolate both lower confined aquifer (A2) and bulk aquifer properties.

### Geochemical Modeling

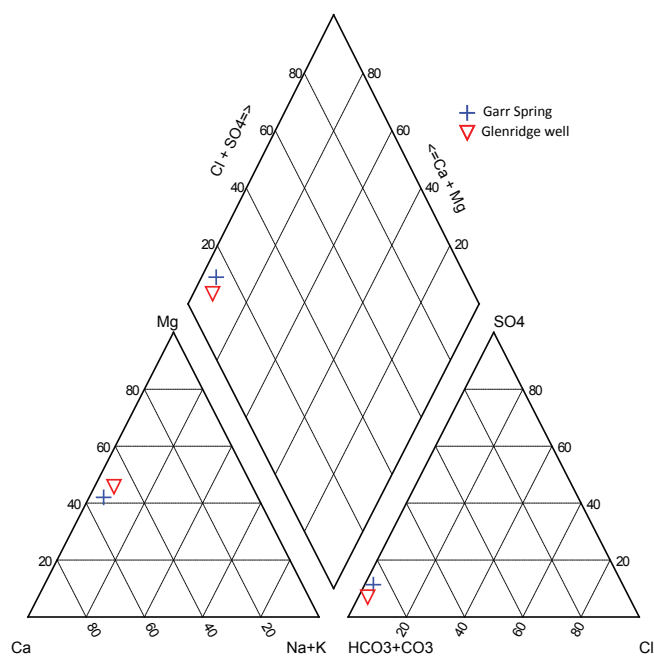
I used PHREEQC software to conduct a very basic mixing model to determine the various phases produced when the Garr Spring water was mixed with the Glenridge well water. I then compared the results of the model to the analysis of samples collected on 3/24/2014. See appendix H for the details of this model.



## RESULTS AND DISCUSSION

### Background Chemistry

Garr Spring and Glenridge well water have very similar geochemistry (figure 11; table 3). Both are calcium-magnesium-bicarbonate water. Geochemical modeling does not suggest that mixing of Garr Spring water with Cache Valley principal aquifer (Glenridge well) water would cause mobilization of nitrate. However, the mixing model does not account for material of other phases (liquid or solid) in the aquifer, which may change phase if exposed to this water. Appendices I and J have the complete results of laboratory analyses.



**Figure 11.** Piper (trilinear) diagram showing Garr Spring and Glenridge well water chemistry.

### Aquifer Properties

Based on the interpolated potentiometric surface of the principal aquifer (figure 12), the horizontal hydrologic gradient near the Glenridge well is 0.004. The mean horizontal hydraulic gradient direction in the area of Millville is down to the west (278 degrees from north) near the well (figure 12).

The aquifers vary in thickness depending on location. Based on the cross section (figures 13 and 14) and the Glenridge well driller's log, the aquifer is about 100 feet thick near the Glenridge well. Most of the clay within 0.25 mile (0.4 km) of the Glenridge well is greater than 80-feet (24m) thick (figure 13). If the interpretation of driller's logs is correct, then the aquifer in the region of Millville coincides well with the conceptual model presented by Robinson (1999) (figure 7). Clay thickness calculations (figure 13) substantiate the re-

charge areas designated in the map from Anderson and others (1994). However, the clay layers thin substantially upgradient (east) of the Glenridge well, where Provo-level Bonneville lake gravels are predominant at the surface (Evans and others, 1996). There may be some thin, intermittent clay layers east of the Glenridge well, allowing for the flow observed at Skinner and Knoll springs (figure 13 and figure 14).

Based on aquifer test analyses (figure 15), the transmissivity of the principal aquifer near the Glenridge well is 135,000 ft<sup>2</sup>/day (12,540 m<sup>2</sup>/day), which, based on an aquifer thickness of 100 feet (30.5 m), equals a hydraulic conductivity of 1350 ft/day (411 m/day). The hydraulic conductivity is extremely high and would be an appropriate value for gravel (Heath, 1983). The transmissivity from the test fits well with other values reinterpolated from Inkenbrandt (2010) (figure 16).

Injection test data were not analyzed due to disruption of the water-level data during injection (figure 17). During the injection portion of the test, the water level in the well dropped about six feet immediately after injection began and stayed at the depressed level until injection stopped (figure 17). This depressed water level is likely due to jetting of water from the injection column (pipe) into the well. There is a significant change in fluid velocity from the port of the injection column, which can lead to a Venturi (Bernoulli) effect in the well casing.

A cone of depression or recharge mound created by a well in a high transmissivity aquifer will exhibit small amounts of potentiometric surface change over a very large area. Based on the transmissivity from the aquifer test using the Glenridge well, injection modeling displays a wide swath of influence. The maximum increase would be about 0.3 feet near the well after an injection cycle and that increase would be negligible after a post-injection pumping cycle. The maximum rate possible for the current injection column is 300 gallons per minute. Based on water rights, Millville would inject from the beginning of October to the end of May the following year (181 days).

Fining of the aquifer material to the west, change of the vertical hydraulic gradient from recharge to discharge zones, and a decrease in horizontal hydraulic gradient all account for lower hydraulic conductivities to the west of Millville (Inkenbrandt, 2010).

### Nitrate Source(s)

The Utah State University Water Laboratory tests detected, but did not measure, small amounts of DEET and pharmaceuticals in water from the Glenridge well, indicating that septic systems are likely contributing contaminants to the principal aquifer system near the well.

Nitrate values interpolated in the area (figure 18; table 4) indicate that an extensive nitrate plume covers most of the central

Station	Date-time	Aluminum	Ammonia as N	Antimony	Arsenic	Barium	Beryllium	Bicarbonate	Boron	Cadmium	Calcium	Carbon Dioxide	Carbonate	Chloride	Chromium	Copper	Cyanide	Fluoride	Hardness	Hydroxide	Iron	Lead	Magnesium
Hancey Well	2/28/14 12:50 PM																						
	3/19/14 4:00 PM																						
Garr Spring	2/28/14 9:15 AM		<0.045	<3	<1	<0.1	<1		0.20						9.7	2.19	<0.01	0.09				0.12	
Glenridge Well	2/28/14 11:15 AM		<0.045	<3	1.09	0.22	<1		0.13						15.9	1.63	<0.01	0.11				0.48	
	3/19/14 2:05 PM		<0.045																				
	3/19/14 10:15 PM																						
	3/19/14 4:30 PM																						
	3/20/14 8:50 AM																						
	3/24/14 10:15 AM	<10	<0.045		<1	0.13		302	<30	<0.1	60.9	10	0	11.7	<2	1.402			274.1	0	<0.02	0.26	31.1
USU WELL	2/28/14 2:06 PM																						
	3/31/14 11:55 AM	<10		<1	<0.1				<30	<0.1	66				<2	<1					0.62	<0.1	25.4

[illegible]

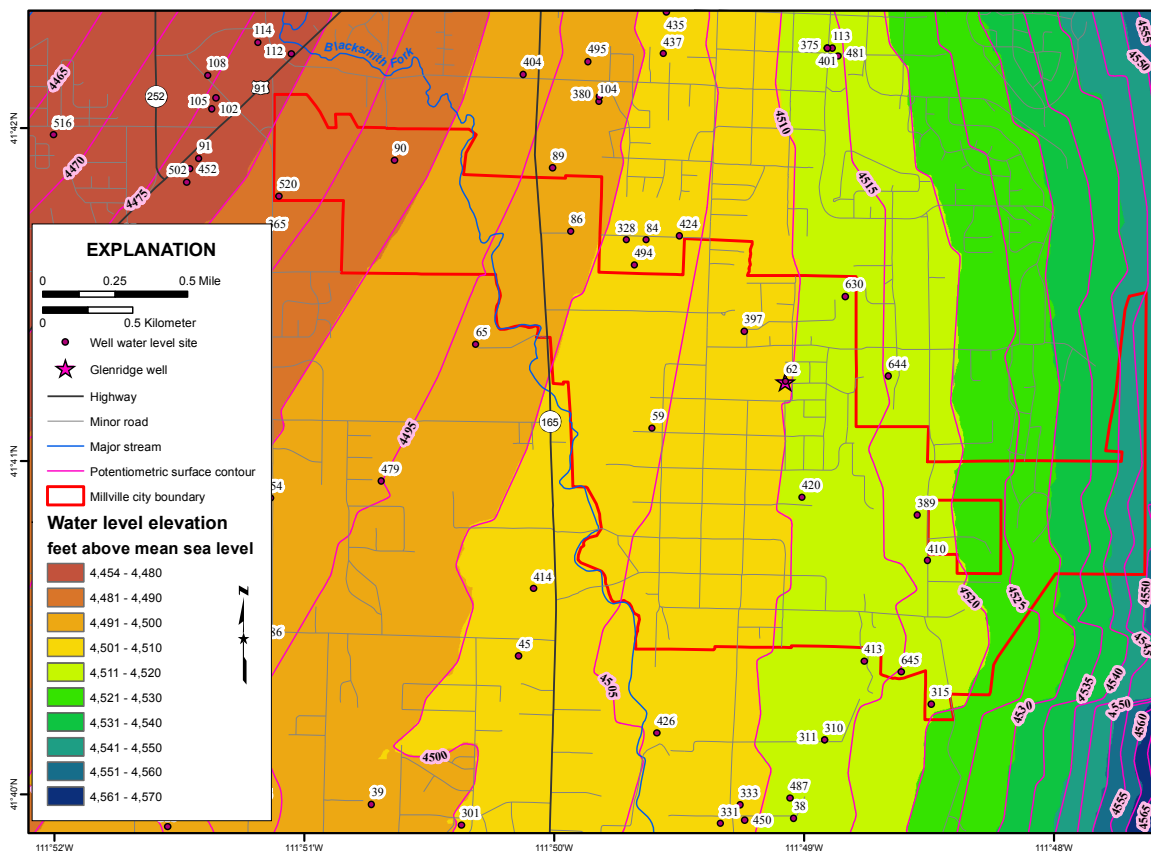


Figure 12. Potentiometric surface contours created using cokriging interpolation of groundwater levels and elevation data in the Millville area. Numbers adjacent to sample sites are sample numbers listed in appendix D.

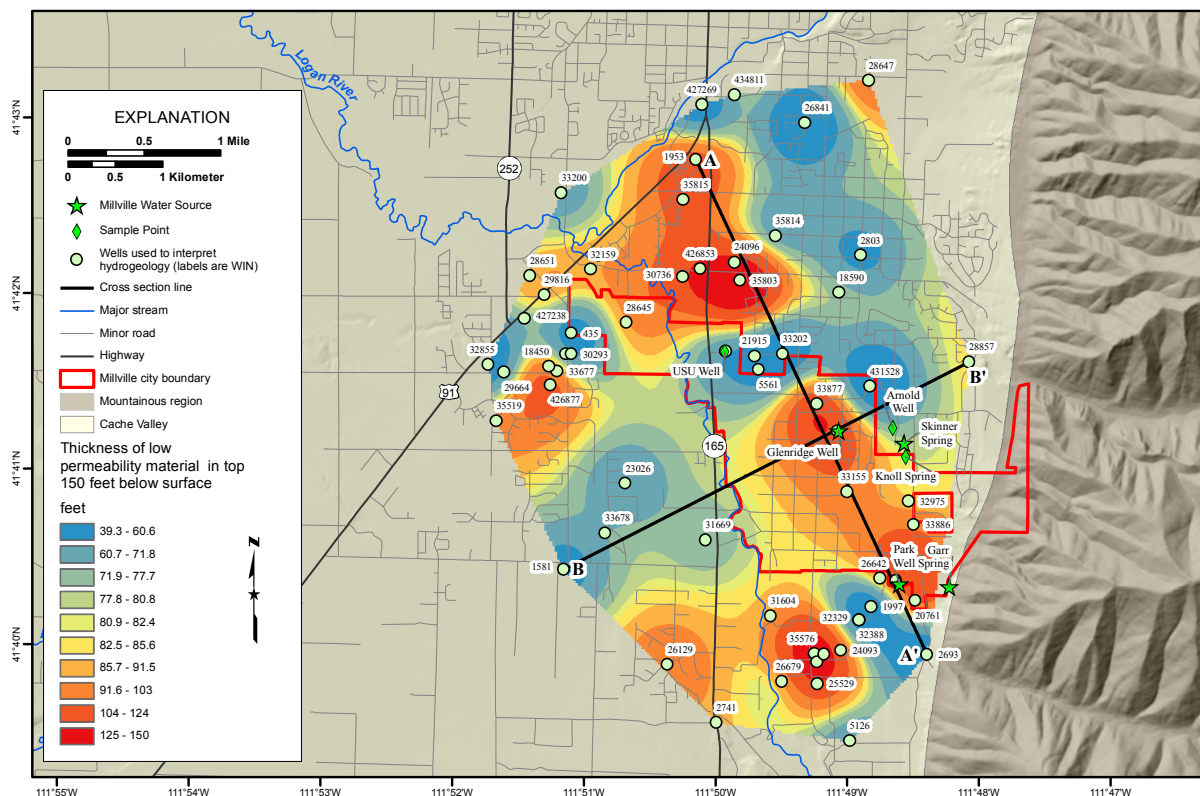
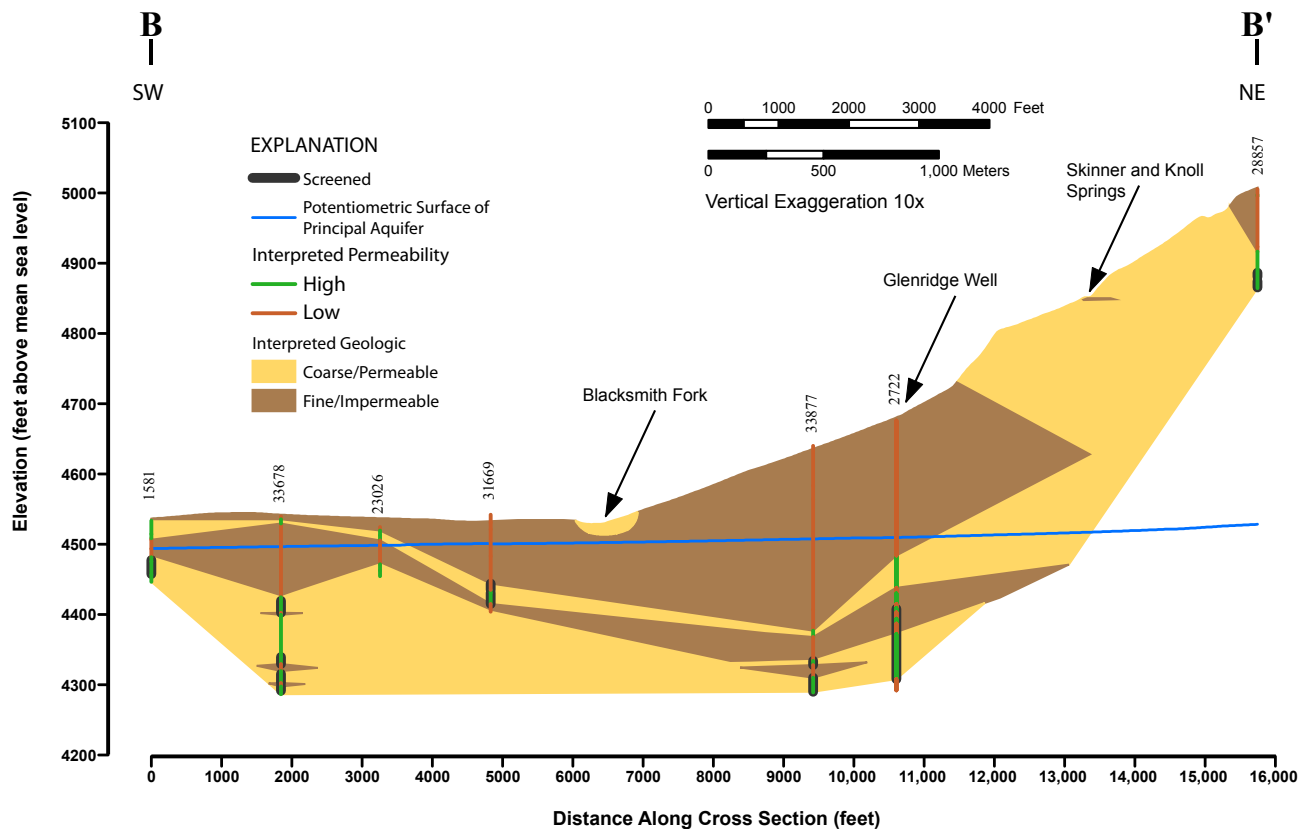
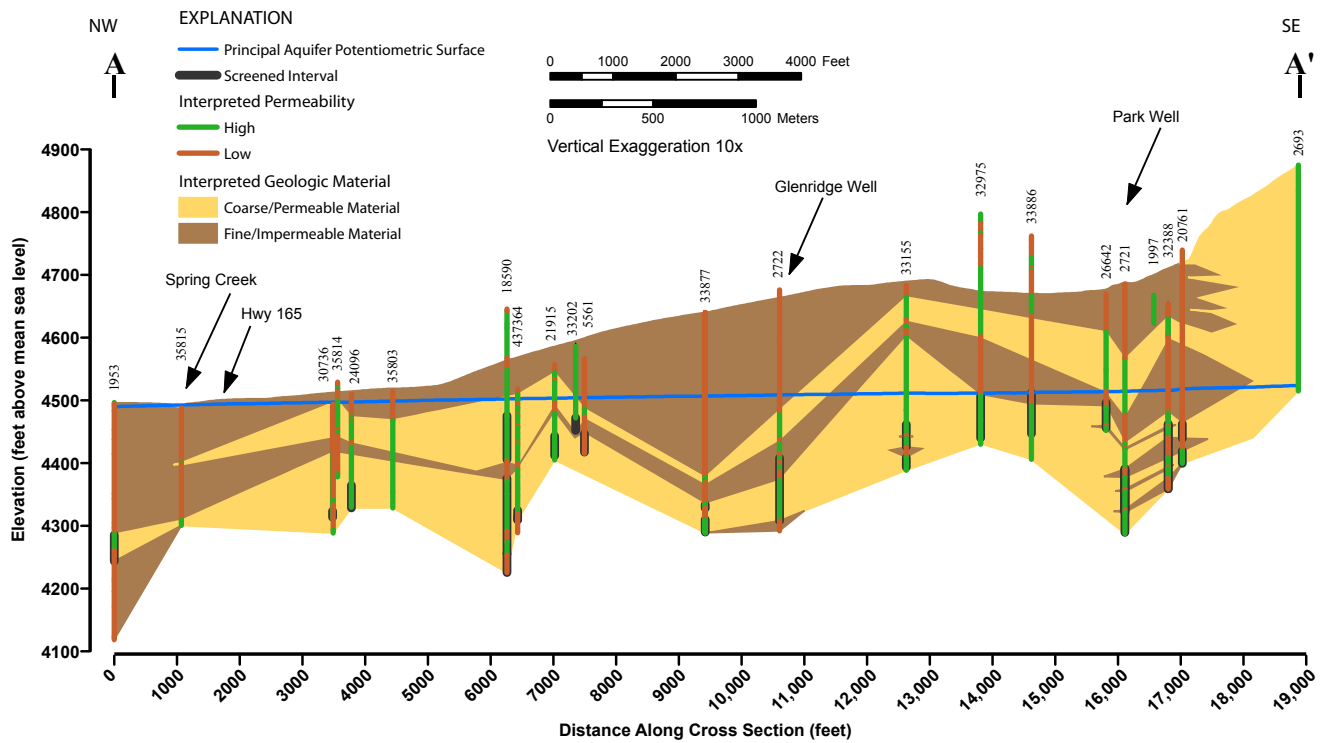


Figure 13. Thickness and distribution of low permeability units and cross section locations. Wells are labeled by their Utah Division of Water Rights well identification number (WIN). See figure 14 for cross sections.



**Figure 14.** Geologic cross sections of Millville area. See figure 13 for location of sections. See appendix E for a summary of the well logs used to make these sections. Numbers above wells indicate the Utah Division of Water Rights well identification number (WIN).



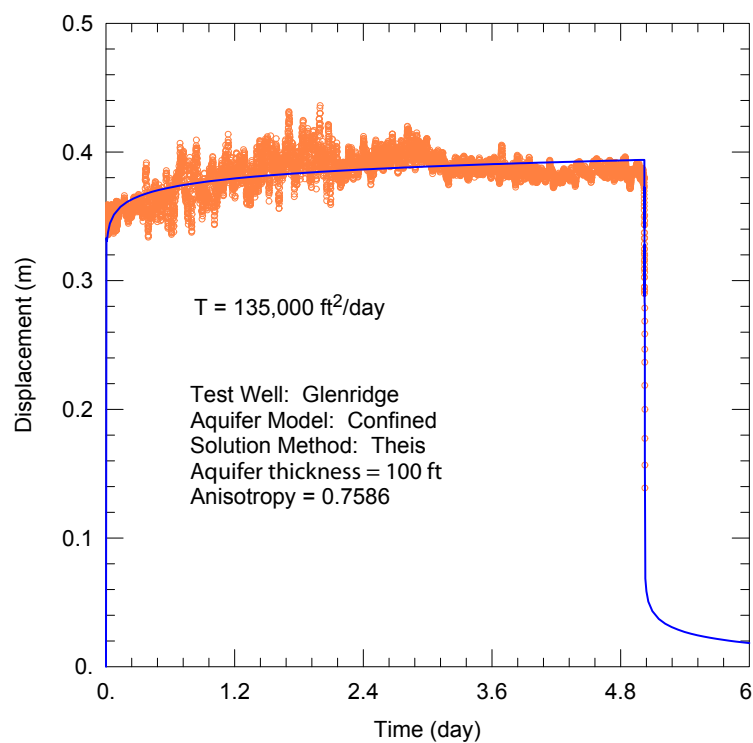


Figure 15. Aquifer test analysis of post-injection pumping data.

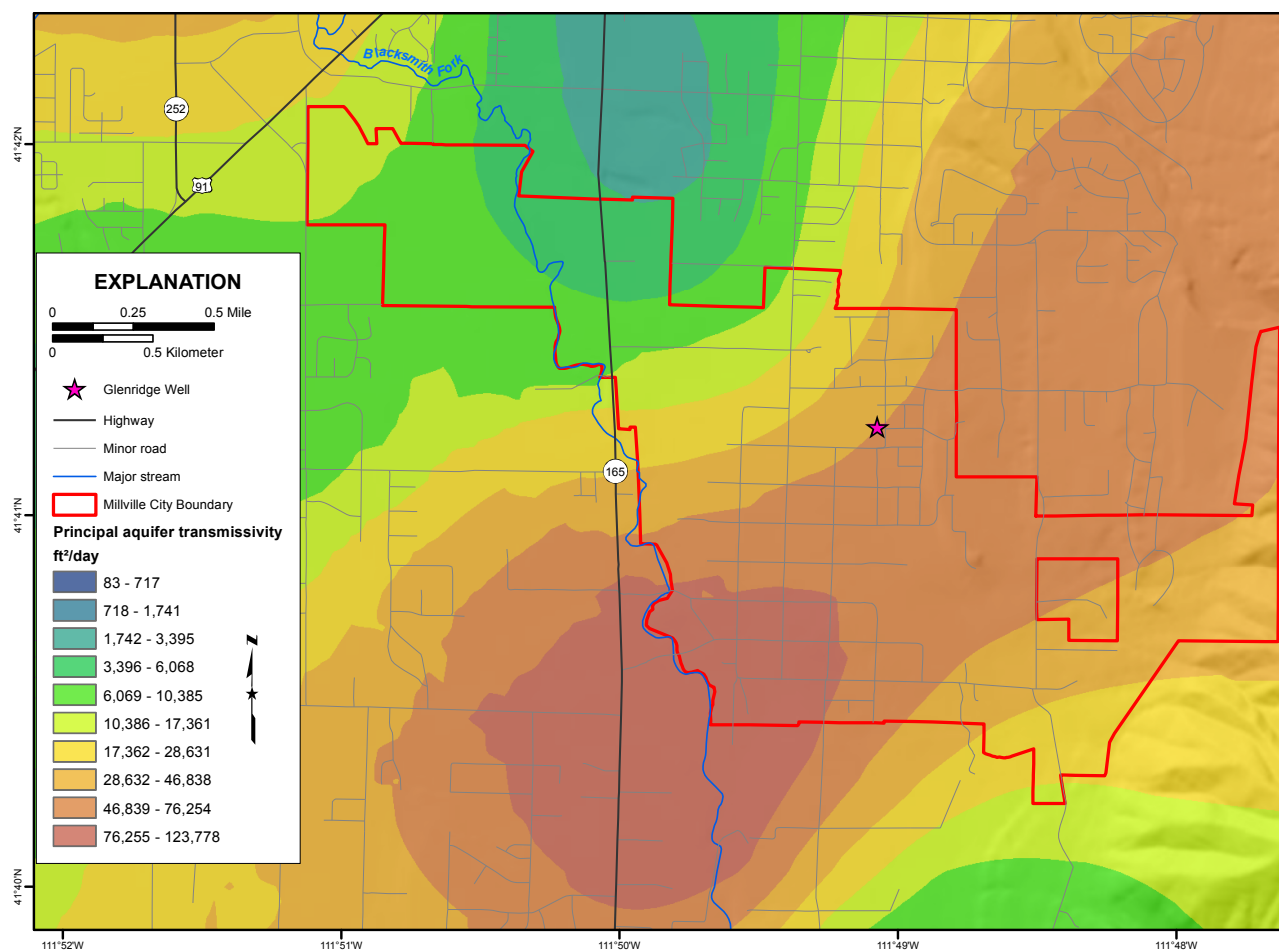


Figure 16. Transmissivity of the Cache Valley principal aquifer modified from Inkenbrandt (2010).

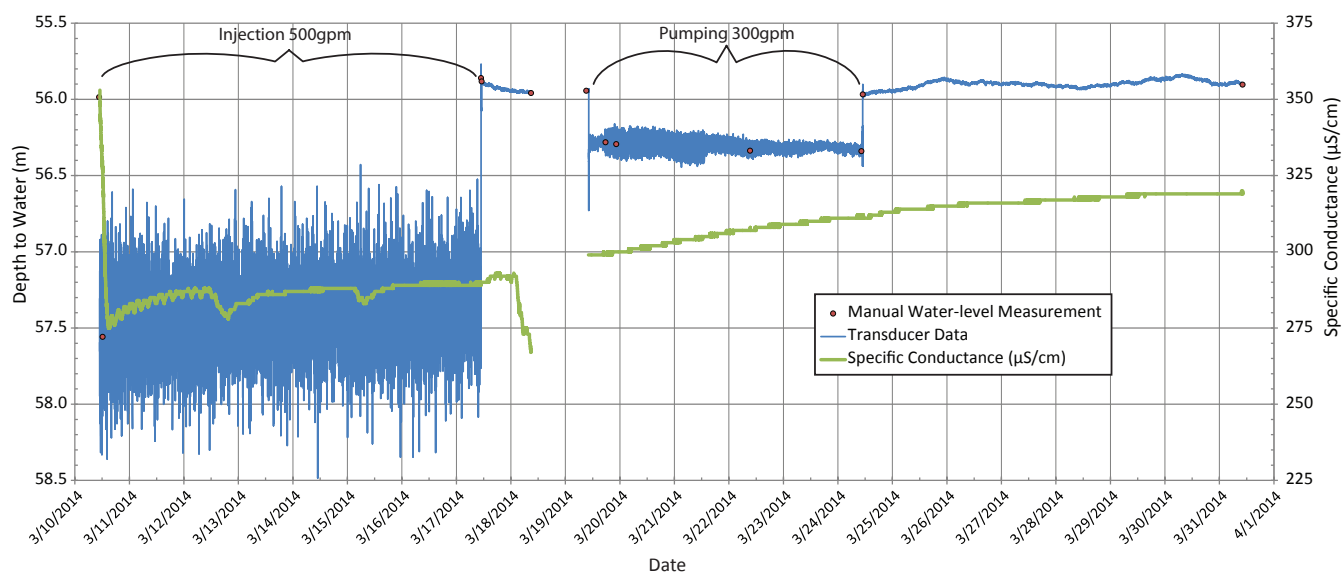


Figure 17. Hydrograph of injection and pumping test of the Glenridge well.

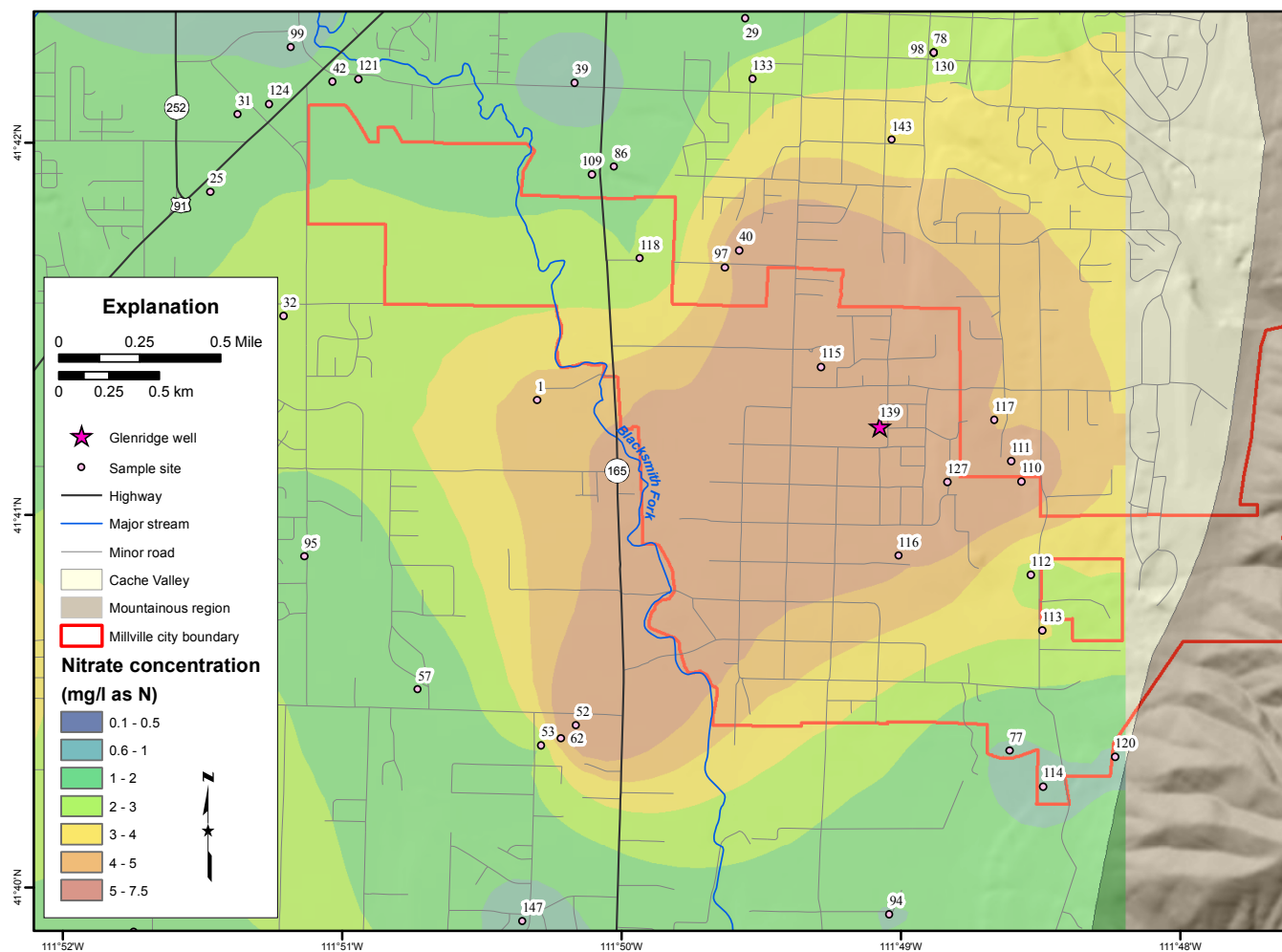


Figure 18. Nitrate concentrations in the Millville area. Numbers adjacent to sample sites are the nitrate sample numbers listed in table 4.

Table 4. Nitrate samples interpolated and examined in this study. See figure 18 for the interpolated values.

Map ID	Longitude	Latitude	Sample Date	Concentration (mg/L NO3-N)	Data Source	Site Name
1	-111.8386	41.6887	11/9/2004	5.2	UGS	CVSS13
2	-111.8532	41.7505	12/3/1997	0.86	UGS	Quayle, James W.
3	-111.8966	41.6541	12/3/1997	0.1	UGS	USU
4	-111.8794	41.7254	11/20/1997	0.43	UGS	Potter, Charles
5	-111.8838	41.7087	11/20/1997	1.46	UGS	Thompson, Leslie
6	-111.8998	41.6545	12/2/1997	0.1	UGS	USU #2
7	-111.9006	41.6904	11/20/1997	0.56	UGS	Jensen, Alvin
8	-111.8955	41.6999	11/20/1997	0.47	UGS	Israelsen, Clark
9	-111.9000	41.6419	12/2/1997	0.1	UGS	Leishman, Kendall
10	-111.8795	41.6469	12/2/1997	4.09	UGS	Miller, Richard L.
11	-111.8978	41.6666	12/2/1997	0.29	UGS	Anderson, A. James
12	-111.8967	41.7217	11/20/1997	0.1	UGS	Thalman, Richard
13	-111.8934	41.6782	11/20/1997	0.46	UGS	Skidmore, Kimberly
14	-111.8902	41.7367	12/2/1997	2.54	UGS	Eliason Packing Co.
15	-111.8913	41.6834	12/16/1997	0.75	UGS	Olsen, Kent L.
16	-111.8894	41.6708	12/16/1997	0.18	UGS	Bridges, Seldon
17	-111.8869	41.6358	12/16/1997	0.1	UGS	Austin, Richard
18	-111.8897	41.6686	12/3/1997	0.61	UGS	Utah State University
19	-111.8726	41.6921	11/19/1997	0.75	UGS	Kunsman, Lisa
20	-111.8649	41.7400	11/20/1997	0.5	UGS	Thalman, Robert
21	-111.8060	41.7213	11/19/1997	3.58	UGS	Andrews, Ronald
22	-111.8604	41.7449	11/20/1997	0.4	UGS	Bodrero, Darrell
23	-111.8155	41.7204	11/19/1997	0.1	UGS	Rounds, Arlyn
24	-111.8642	41.6926	11/19/1997	1.4	UGS	Hansen, Kay D.
25	-111.8583	41.6979	11/19/1997	1.47	UGS	Russell, Bert
26	-111.8634	41.6752	11/19/1997	1.38	UGS	Floyd, W.D.
27	-111.8734	41.7226	11/19/1997	2.32	UGS	Jensen, Robert L.
28	-111.8730	41.7017	11/19/1997	1.41	UGS	Isaacson, Merl
29	-111.8264	41.7059	11/19/1997	0.57	UGS	Gustaveson, Rex
30	-111.8605	41.7317	11/20/1997	0.1	UGS	Clark, Darala & Merrill, Glacus
31	-111.8567	41.7013	11/19/1997	1.58	UGS	Hyclone
32	-111.8538	41.6923	11/19/1997	1.88	UGS	Smith, Claine
33	-111.8674	41.7390	11/20/1997	0.1	UGS	Holographic Products
34	-111.8744	41.6852	11/19/1997	0.7	UGS	Hansen, Hal
35	-111.8371	41.7102	11/19/1997	3.14	UGS	Weston, Todd G.
36	-111.8459	41.6621	11/19/1997	2.14	UGS	Peterson, Steve and Cindy
37	-111.8352	41.6476	11/18/1997	0.51	UGS	Larsen, Kent
38	-111.8639	41.6556	11/19/1997	0.26	UGS	Miller, E.A. - Corp.
39	-111.8366	41.7029	11/19/1997	0.1	UGS	Alder, Seth L.
40	-111.8267	41.6955	11/19/1997	9.71	UGS	Olsen, David
41	-111.8736	41.6932	11/19/1997	0.58	UGS	Jenson, Edwin
42	-111.8510	41.7029	11/19/1997	1.54	UGS	Zollinger, LA
43	-111.8206	41.7158	11/19/1997	0.83	UGS	Smith, Arthur D.
44	-111.8630	41.6930	11/19/1997	1.16	UGS	Isaacson, Merle
45	-111.8624	41.6647	11/19/1997	1.25	UGS	Wright, Steven
46	-111.8716	41.6746	12/16/1997	4.03	UGS	Zollinger, Sid
47	-111.8970	41.6245	5/13/1998	0.27	UGS	Nielsen
48	-111.8486	41.6433	3/9/1998	0.352	UGS	well 1, hyrum
49	-111.8567	41.6475	3/9/1998	3.46	UGS	well 2, hyrum
50	-111.8686	41.6517	3/9/1998	5.23	UGS	well 4, hyrum
51	-111.8792	41.6522	3/9/1998	3.54	UGS	well 5, hyrum
52	-111.8361	41.6742	9/1/2004	9	UDAF	1488
53	-111.8382	41.6732	9/6/2000	1.5	UDAF	1489
54	-111.8927	41.6304	9/1/2004	3.2	UDAF	1490
55	-111.8911	41.6833	9/13/2000	1.2	UDAF	1492
56	-111.8912	41.6850	9/13/2000	1.5	UDAF	1493
57	-111.8456	41.6757	10/10/2000	0.9	UDAF	1494
58	-111.8617	41.6747	8/1/2002	2	UDAF	1501
59	-111.8169	41.6461	8/1/2002	0.6	UDAF	1502
60	-111.8168	41.6458	8/1/2002	0.8	UDAF	1503
61	-111.8905	41.6850	9/1/2004	1	UDAF	1509

Table 4. Continued

Map ID	Longitude	Latitude	Sample Date	Concentration (mg/L NO3-N)	Data Source	Site Name
62	-111.8370	41.6736	9/1/2004	5.8	UDAF	1511
63	-111.8932	41.6297	9/1/2004	3.5	UDAF	1514
64	-111.8903	41.6306	9/1/2004	2	UDAF	1523
65	-111.8553	41.7336	9/6/2000	0.5	UDAF	1569
66	-111.8574	41.7236	6/27/2001	3.8	UDAF	1572
67	-111.8621	41.7106	9/1/2004	0.1	UDAF	1590
68	-111.8853	41.7192	9/6/2000	0.8	UDAF	2823
69	-111.8935	41.6985	9/6/2000	0.1	UDAF	2824
70	-111.8602	41.7279	9/13/2006	0.8608	UDAF	3162
71	-111.8734	41.6931	9/13/2006	0.8022	UDAF	3165
72	-111.8933	41.7185	9/13/2006	0.4634	UDAF	3166
73	-111.8191	41.6427	6/20/1988	1	STORET_Leg	300102
74	-111.8187	41.6413	6/20/1982	3.2	STORET_Leg	300103
75	-111.8336	41.6341	6/20/1988	1	STORET_Leg	300803
76	-111.8319	41.6272	6/20/1988	0.5	STORET_Leg	300804
77	-111.8102	41.6732	6/20/1978	0.6	STORET_Leg	301203
78	-111.8151	41.7045	6/20/1988	2.6	STORET_Leg	301702
79	-111.8261	41.7235	6/20/1985	1.7	STORET_Leg	301901
80	-111.8186	41.7233	6/20/1988	0.4	STORET_Leg	301904
81	-111.8972	41.6588	6/20/1988	0.2	STORET_Leg	309001
82	-111.8897	41.6981	12/2/1980	0.146835	STORET	4904980
83	-111.8997	41.7394	9/24/1991	0.249619	STORET	4905070
84	-111.8686	41.7364	2/7/1979	0.2259	STORET	4905110
85	-111.8694	41.7367	12/2/1980	0.1	STORET	4905120
86	-111.8342	41.6992	6/6/1978	0.4518	STORET	4905410
87	-111.8803	41.6519	7/20/1988	2.085057	STORET	4905520
88	-111.8683	41.6558	7/9/1985	0.137799	STORET	4905540
89	-111.8905	41.6269	4/18/1968	0.1	USGS_NWIS	(A-10- 1) 6ccc- 1
90	-111.8333	41.6338	6/17/1968	0.203	USGS_NWIS	(A-10- 1) 4daa- 1
91	-111.8605	41.6438	3/28/1968	0.136	USGS_NWIS	(A-11- 1)32dc- 1
92	-111.8924	41.6466	3/20/1968	0.339	USGS_NWIS	(B-11- 1)36dad- 1
93	-111.8272	41.6566	7/4/1968	0.361	USGS_NWIS	(A-11- 1)27cdc- 1
94	-111.8173	41.6658	10/6/1998	0.826	USGS_NWIS	(A-11- 1)27adb- 1
95	-111.8524	41.6816	7/4/1968	0.745	USGS_NWIS	(A-11- 1)20ada- 1
96	-111.8733	41.6858	8/31/1962	0.339	USGS_NWIS	(A-11- 1)18ddd- 1
97	-111.8275	41.6947	8/17/2011	3.85	USGS_NWIS	(A-11- 1)15bdb- 1 S29
98	-111.8152	41.7044	8/31/1966	2.48	USGS_NWIS	(A-11- 1)10dad- 1
99	-111.8536	41.7044	4/4/1961	0.361	USGS_NWIS	(A-11- 1) 8dda- 3
100	-111.8416	41.7102	7/20/1960	1.69	USGS_NWIS	(A-11- 1) 9acb- 2
101	-111.8258	41.7244	3/12/1954	1.6	USGS_NWIS	(A-11- 1) 3bda- 1
102	-111.8202	41.7255	3/12/1954	2.17	USGS_NWIS	(A-11- 1) 3aca- 1
103	-111.8286	41.7313	2/1/1963	0.271	USGS_NWIS	(A-12- 1)34cca- 1
104	-111.8772	41.7358	8/31/1962	0.429	USGS_NWIS	(A-12- 1)31dab- 1
105	-111.8705	41.7360	2/6/1963	0.294	USGS_NWIS	(A-12- 1)32cbb- 1
106	-111.8127	41.7380	3/14/1963	0.203	USGS_NWIS	(A-12- 1)35bcc- 1
107	-111.8194	41.7447	6/15/1962	0.136	USGS_NWIS	(A-12- 1)27dcd- 1
108	-111.8274	41.7497	1/3/1964	0.181	USGS_NWIS	(A-12- 1)27cab- 1
109	-111.8355	41.6988	5/8/1968	3.16	USGS_NWIS	(A-11- 1)15bbc-S1
110	-111.8097	41.6853	6/20/2012	6.89	Millville City	Knowles Springs
111	-111.8103	41.6862	6/20/2012	6.03	Millville City	Mathews Spring
112	-111.8091	41.6811	6/20/2012	1.91	Millville City	Owen Hancey Well
113	-111.8084	41.6786	6/20/2012	3.48	Millville City	K. Hancey Well
114	-111.8082	41.6716	6/20/2012	0.73	Millville City	Postma Well
115	-111.8217	41.6903	6/20/2012	6.83	Millville City	LEGRAND MATHEWS
116	-111.8170	41.6819	6/20/2009	5.43	Millville City	Cox Well
117	-111.8113	41.6880	3/21/2014	4.07	UGS	Arnold Well
118	-111.8326	41.6951	2/28/2014	1.08	UGS	<Null>
119	-111.8097	41.7433	9/18/2012	0.6	SDWIS	NATURAL RESOURCES BUILDING WELL
120	-111.8039	41.6730	8/24/2012	0.778	SDWIS	GARR SPRING
121	-111.8495	41.7030	8/23/2012	1.24	SDWIS	ZOLLINGER WELL
122	-111.8226	41.6444	7/31/2012	5.4	SDWIS	NUMBER ONE SPRING



Table 4. Continued

Map ID	Longitude	Latitude	Sample Date	Concentration (mg/L NO3-N)	Data Source	Site Name
123	-111.8264	41.7243	12/13/2005	2.01	SDWIS	LOWER WELL
124	-111.8548	41.7018	11/30/1998	1.52	SDWIS	ARTESIAN WELL
125	-111.8972	41.6588	2/24/1998	0.4	SDWIS	THATCHER WELL
126	-111.9014	41.6587	9/8/1994	0.1	SDWIS	NEW WELL
127	-111.8141	41.6852	8/23/2012	6.03	SDWIS	MATTHEWS SPRING
128	-111.8133	41.7382	1/11/2012	0.4	SDWIS	CROCKETT AVE#1 WELL
130	-111.8150	41.7043	8/15/2007	2.71	SDWIS	DALES WELL 100 E 200 S
131	-111.8205	41.7247	12/13/2005	1.86	SDWIS	UPPER WELL
132	-111.8335	41.6339	11/1/2012	0.7	SDWIS	WELL #1
133	-111.8260	41.7032	9/27/2012	3.6	SDWIS	ALDER-WEST WELL
134	-111.8072	41.7439	9/18/2012	0.5	SDWIS	INDUSTRIAL SCIENCE WELL
136	-111.8289	41.7306	8/15/2012	0.5	SDWIS	200 E CENTER WELL
138	-111.8997	41.6598	11/14/1991	0.1	SDWIS	OLD FARM N WELL
139	-111.8182	41.6877	8/24/2012	7.81	SDWIS	GLENRIDGE WELL
140	-111.8332	41.6600	7/31/2012	0.5	SDWIS	4000 SOUTH MAIN WELL
141	-111.8189	41.7229	6/29/2012	2.4	SDWIS	MUNICIPAL WELL
142	-111.8196	41.7444	1/11/2012	0.4	SDWIS	700 N 600 E WELL
143	-111.8176	41.7005	9/30/2008	2.84	SDWIS	400 S PROVIDENCE WELL
144	-111.8319	41.6272	10/14/1988	0.57	SDWIS	WELL #2 DISCONNECTED
145	-111.8478	41.7194	8/15/2012	0.4	SDWIS	WILLOW PARK WELL
146	-111.8188	41.6402	9/4/1990	0.51	SDWIS	THOMAS IRR WELL
147	-111.8392	41.6654	7/31/2012	0.4	SDWIS	NELSON WELL
148	-111.8279	41.7495	11/7/2002	0.84	SDWIS	1000 N 300 E WELL
150	-111.9347	41.6924	7/21/1948	0.045001	WQP	USGS-414133111560201
152	-111.9111	41.7416	6/30/1948	0.02259	WQP	USGS-414430111543701
155	-111.9180	41.7860	7/12/1948	1.310019	WQP	USGS-414710111550201
158	-111.9791	41.7874	10/27/1949	0.1	WQP	USGS-414715111584201
159	-111.9277	41.8005	7/12/1948	0.090001	WQP	USGS-414802111553701
168	-111.9077	41.8447	10/27/1949	0.1	WQP	USGS-415041111542501
176	-111.9163	41.5930	4/18/1968	2.71004	WQP	USGS-413535111545601
178	-111.8474	41.6110	3/20/1968	2.259	WQP	USGS-413640111504801
180	-111.9436	41.6185	4/18/1968	1.310019	WQP	USGS-413707111563400
182	-111.9327	41.6244	5/9/1968	1.92015	WQP	USGS-413728111555501
186	-111.9338	41.6444	4/18/1968	1.53612	WQP	USGS-413840111555901
192	-111.9211	41.6472	5/27/1968	0.068001	WQP	USGS-413850111551301
194	-111.9183	41.6488	3/20/1968	0.904013	WQP	USGS-413856111550301
196	-111.9313	41.6733	5/27/1968	0.1	WQP	USGS-414024111555001
197	-111.9536	41.6752	5/7/1968	4.518	WQP	USGS-414031111571001
199	-111.8824	41.6941	5/8/1968	0.565008	WQP	USGS-414139111525401
201	-111.9111	41.6994	4/18/1968	0.1	WQP	USGS-414158111543701
202	-111.9619	41.7024	4/17/1968	0.15813	WQP	USGS-414209111574001
204	-112.0069	41.7119	9/4/1959	0.520008	WQP	USGS-414243112002201
211	-111.9358	41.7463	4/18/1968	0.02259	WQP	USGS-414447111560601
213	-111.8388	41.7555	7/13/1960	0.06777	WQP	USGS-414520111501701
215	-111.8444	41.7569	7/13/1960	0.045001	WQP	USGS-414525111503705
217	-111.8061	41.7588	3/31/1955	1.36002	WQP	USGS-414532111481901
227	-111.8352	41.7747	4/17/1968	3.610053	WQP	USGS-414629111500401
229	-111.8327	41.7760	9/30/1965	1.67166	WQP	USGS-414634111495501
231	-111.8894	41.7794	4/18/1968	0.045001	WQP	USGS-414646111531901
233	-111.8536	41.7827	5/27/1959	0.99396	WQP	USGS-414658111511001
237	-111.9788	41.7874	8/21/1950	0.06777	WQP	USGS-414715111584101
243	-111.8338	41.7927	4/11/1963	0.76806	WQP	USGS-414734111495901
245	-111.9047	41.7952	12/10/1957	0.15813	WQP	USGS-414743111541401
251	-111.8311	41.8066	8/19/1966	4.290063	WQP	USGS-414824111494901
253	-111.9286	41.8088	4/18/1968	0.1	WQP	USGS-414832111554001
254	-111.8486	41.8152	5/8/1968	3.840057	WQP	USGS-414855111505201
256	-111.8574	41.8247	5/8/1968	5.1957	WQP	USGS-414929111512401
258	-111.8647	41.8313	7/20/1960	0.40662	WQP	USGS-414953111515001
268	-111.8349	41.8449	7/9/1968	1.76202	WQP	USGS-415042111500301
270	-111.8883	41.8499	7/9/1957	0.068001	WQP	USGS-415100111531501
314	-111.8686	41.8388	9/2/1970	0.045001	WQP	USGS-415020111520401

Table 4. Continued

Map ID	Longitude	Latitude	Sample Date	Concentration (mg/L NO <sub>3</sub> -N)	Data Source	Site Name
316	-111.9514	41.6474	8/19/1998	3.8403	WQP	USGS-413850111570301
322	-111.8908	41.6781	4/4/2001	0.480007	WQP	UTAHDWQ-4904060
323	-111.8585	41.6748	4/4/2001	2.490037	WQP	UTAHDWQ-4904010
325	-111.8817	41.6925	4/4/2001	1.260019	WQP	UTAHDWQ-4904050
327	-111.8762	41.6776	4/4/2001	14.700217	WQP	UTAHDWQ-4904020
329	-111.8622	41.6897	4/4/2001	3.830057	WQP	UTAHDWQ-4904030

part of Millville. The contamination extends in both the unconfined and confined aquifers in the principal aquifer system, as observed by measured values of nitrate in both springs and deeper wells.

A map of septic tank density of the Millville area (figure 19) indicates a large quantity of septic systems upgradient of Glenridge well. However, based on observed nitrate concentrations of Skinner and Knoll springs, the distribution of confining clay thickness in the area (figure 13), and the extent of the nitrate contamination plume, major contributing sources likely exist upgradient of the springs. While most of the parcels within the town of Providence are connected to the city's sewer system, there are likely some exceptions immediately upgradient of Skinner and Knoll springs where older homes

are present that may not have connected to the Providence system. After carefully reviewing historical aerial photography and land-use maps, I found no major livestock or agricultural activities upgradient of the Glenridge well. However, a small farm is located east of the well, with some llamas and evidence of historic manure piles, that could potentially contribute to nitrate observed in the Glenridge well.

Results of analyses of nitrogen and oxygen isotopes in nitrate from Glenridge well water are ambiguous. The average nitrogen and oxygen isotope concentration ratios were 5.75 and -5.78 ‰, respectively. Many different nitrate sources could produce these isotope concentration ratios (figure 20). However, based on these values, manufactured nitrate is likely not a source.

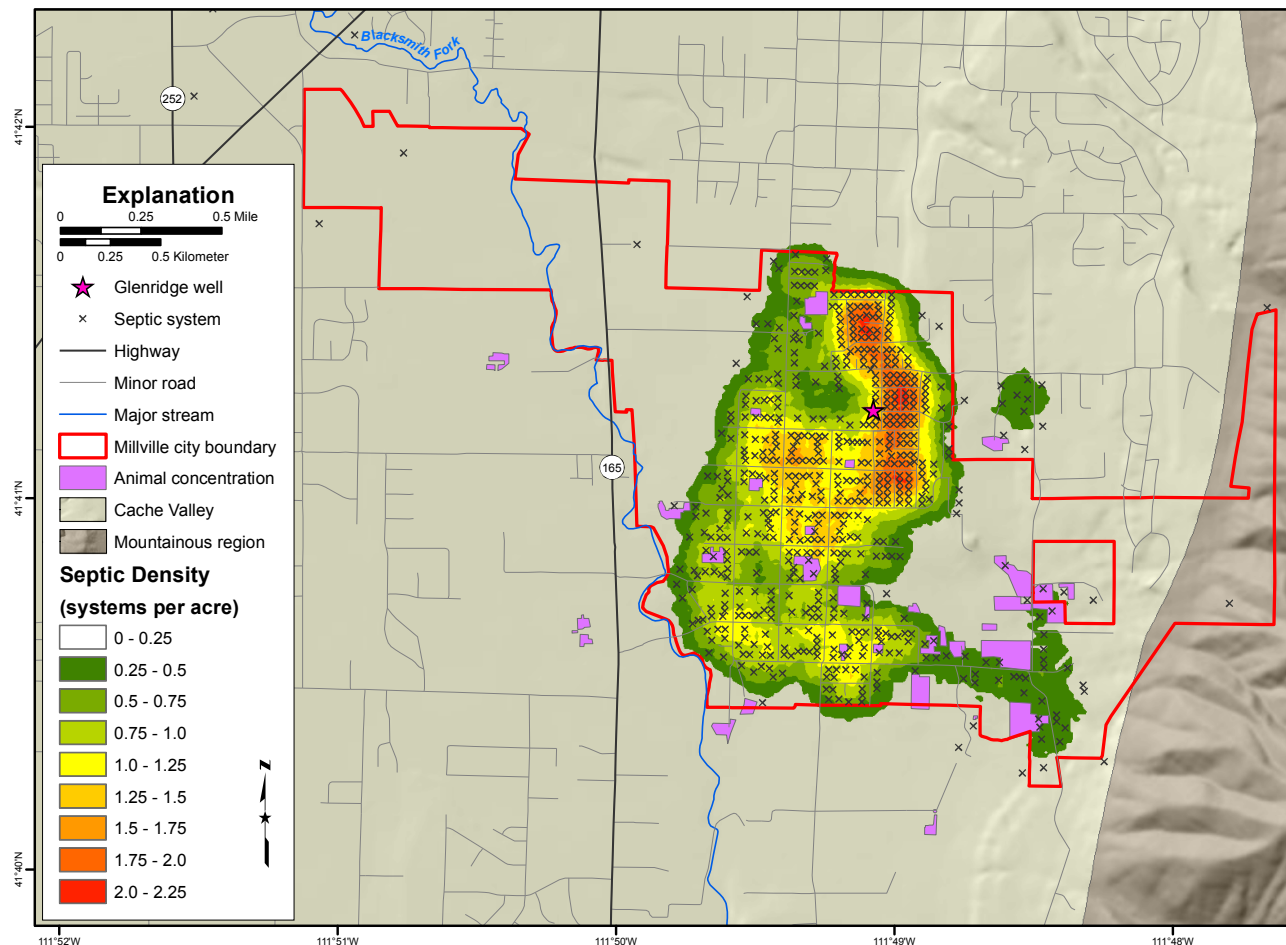


Figure 19. Septic tank locations, septic tank density, and locations of animal concentrations.

Recovery Efficiency

During the initial post-injection pumping, Millville pumped out 42.5% of the total volume of water that was injected. However, the water extracted did not have the same composition of the water that was injected. I used ratios of oxygen and deuterium isotopes (table 5; figure 21) and nitrate concentrations (figure 22; table 4) to determine how much of the pumped water was the injectate.

Isotope and nitrate concentrations show that the water extracted from the Glenridge well was 95% Garr Spring water up to 10.75 hours after post-injection pumping began (figure 22). For the first five days of pumping, nitrate concentration increased in a near-linear fashion at a rate of about 0.5 mg/l for every acre-feet of water pumped. At the end of the initial interval of pumping, on 3/24/2014 at 10:30 a.m., the water extracted from the well was 68% Garr Spring water (figure 23). Nitrate values increased logarithmically during the post-injection pumping period (figure 22a). This is best explained by diffusion of the Garr Spring water with the original aquifer water. Relatively fast-flowing groundwater likely exists near the well, allowing for further mixing of the two waters. The asymptote of the nitrate trend approaches 7 mg/l, which is a relatively high value, but lower than the initial value.

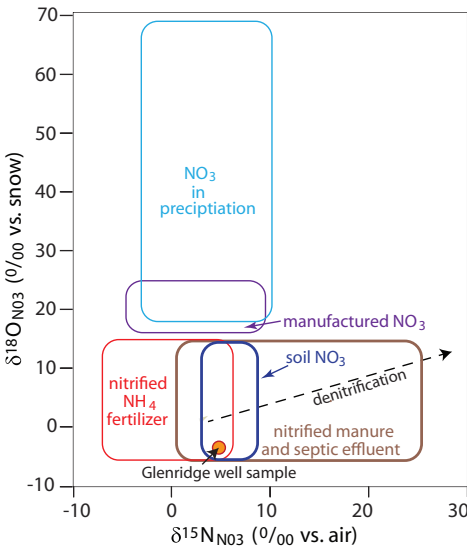


Figure 20. Standard ranges for isotope concentrations of nitrate-15 and oxygen-18 for various nitrate sources (modified from Kendall, 1998). Nitrate from water in the Glenridge well falls within several categories.

Table 5. Results of stable isotope analyses.

Site	Date-Time	d18O (‰)	+/- d18O (‰)	dD (‰)	+/- dD (‰)
USU well	3/31/2014 11:55	-16.95	0.04	-126.0	3.1
Knoll Spring	3/31/2014 10:00	-15.93	0.04	-114.7	3.1
Arnold well	3/21/2014 10:00	-16.46	0.04	-123.6	3.1
Glenridge well	3/19/2014 22:15	-16.88	0.04	-128.1	3.1
Garr Spring	2/28/2014	-16.92	0.04	-127.1	3.1
Glenridge well	3/24/2014	-16.70	0.04	-123.8	3.1
Glenridge well	2/28/2014 11:45	-16.15	0.04	-115.5	3.1
Glenridge well	4/17/2014	-16.50	0.06	-121.8	2.1

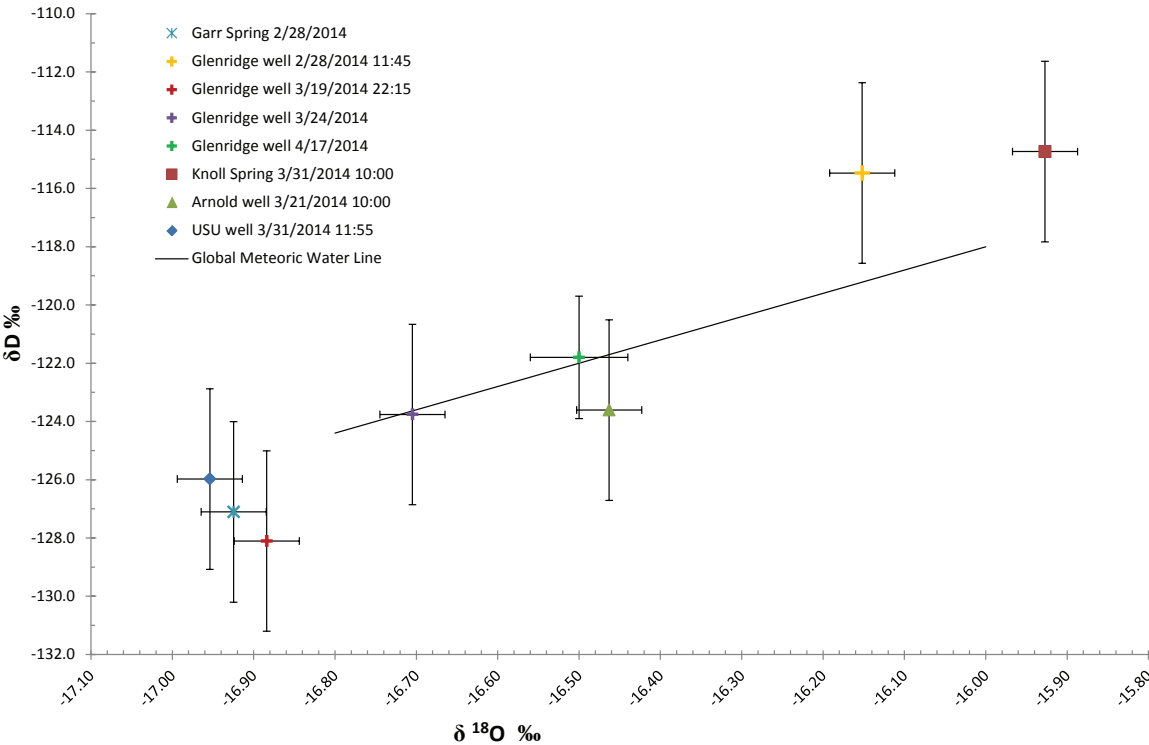
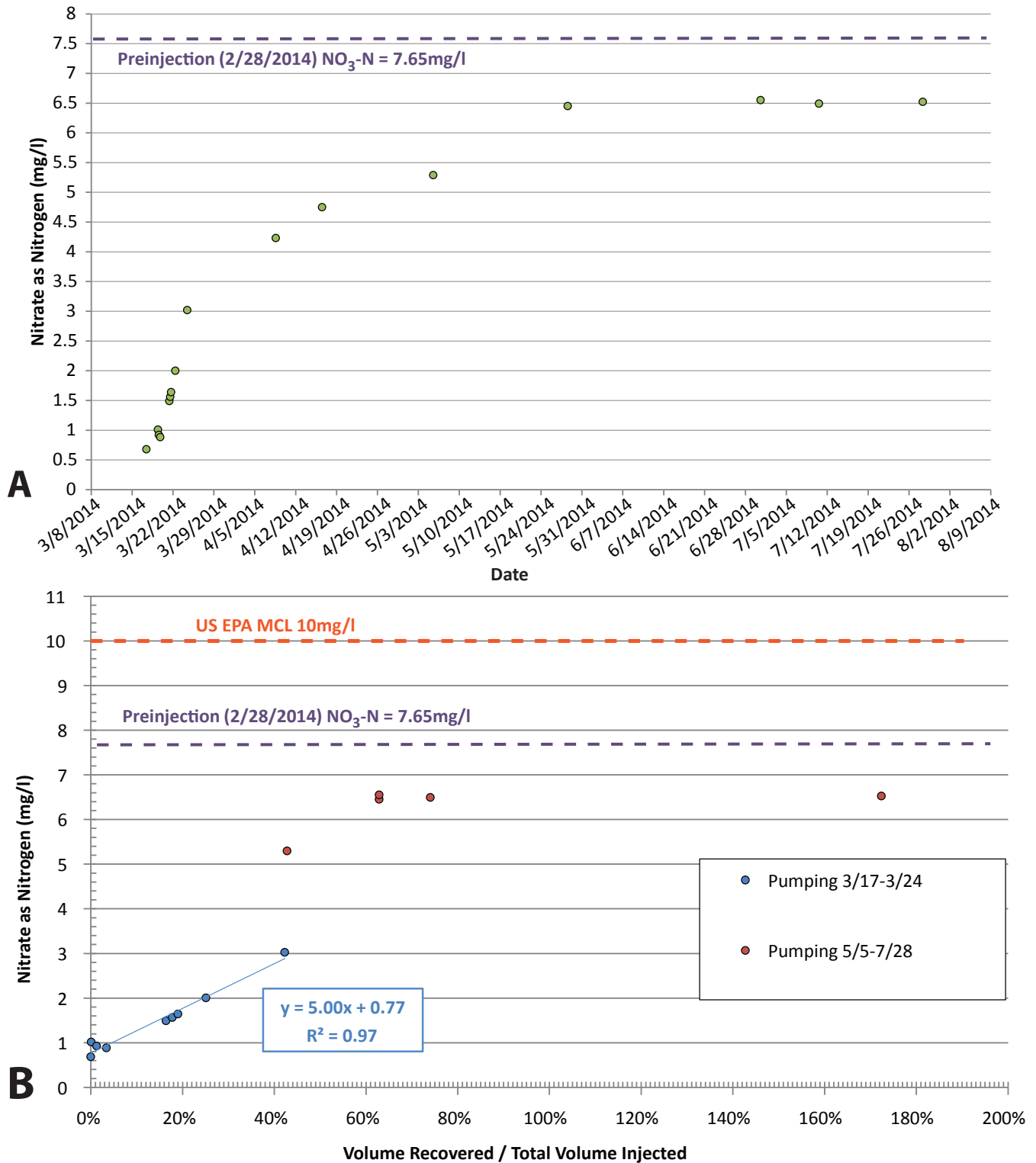
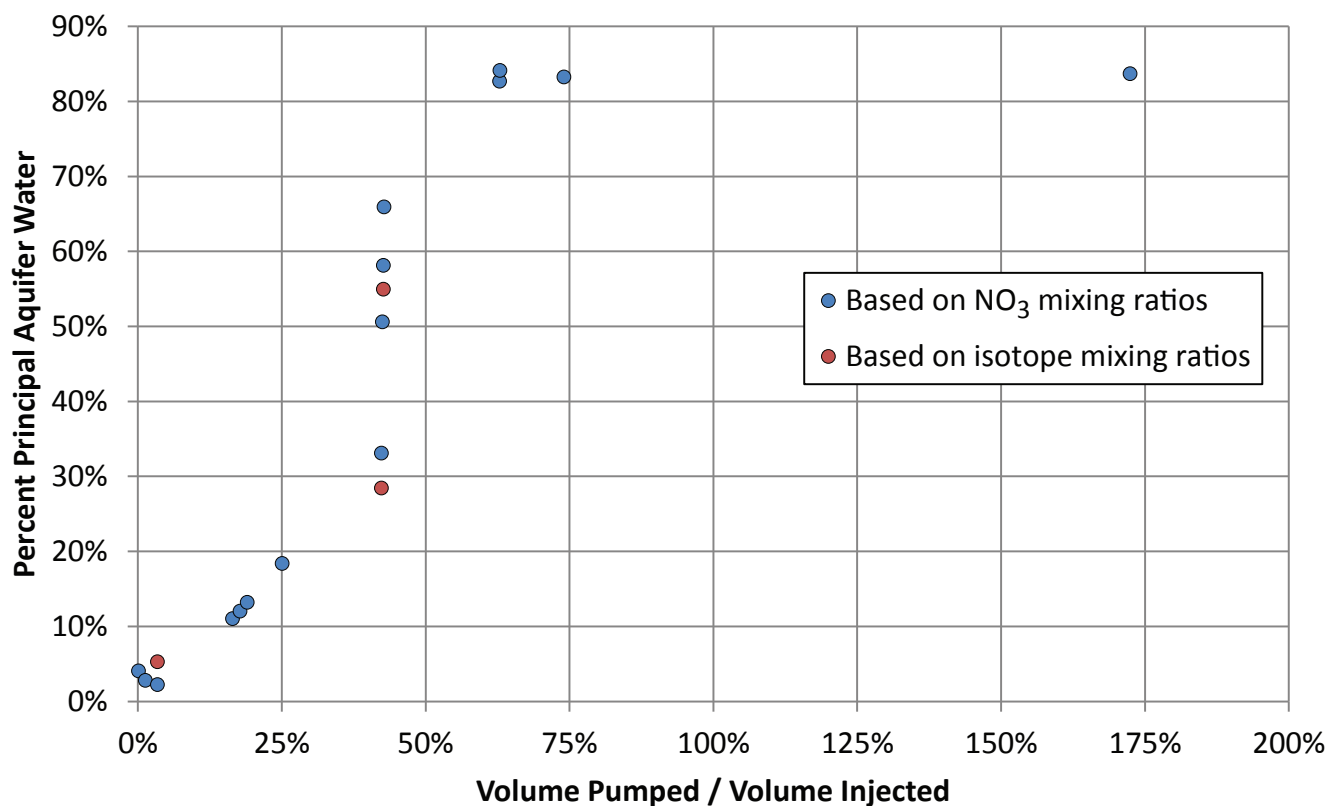


Figure 21. Deuterium and oxygen stable isotope ratios from sources in Millville area.



**Figure 22.** A) Concentration of nitrate in the Glenridge well over time for the duration of post-injection pumping. B) Concentration of nitrate as a function of volume of water pumped from the Glenridge well.





**Figure 23.** Percentage of pumped water that is original to the Cache Valley principal aquifer relative to the amount of water pumped from the Glenridge well.

## SUMMARY

The aquifer system is likely contaminated upgradient (east) of the Glenridge well, especially in areas to the northeast where clay layers are thin. Much of the area of greatest apparent vulnerability is within Providence City limits.

Based on a short-term preliminary injection test, introducing Garr Spring water into Glenridge well does not significantly alter the chemistry of the Cache Valley principal aquifer system and effectively stores Garr Spring water in the aquifer. However, prolonged residence time of Garr Spring water could allow dilution of the spring water with native aquifer water.

## RECOMMENDATIONS

I suggest that Millville locate and remediate the high nitrate sources upgradient of their well. A more thorough analysis of Glenridge well water for common byproducts of septic contamination is recommended before injection occurs. A longer-term injection test is necessary before impact on the aquifer system can be successful. For the long-term test, the chemistry of the solid phase of the geochemical system should be modeled.

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