

DESCRIPTION OF MAP UNITS

Qa ALLUVIAL DEPOSITS (QUATERNARY)—Silt, sand, and gravel in alluvial fans, alluvial slope wash, and stream alluvium.

Qm LANDSLIDE DEBRIS (QUATERNARY)—Includes talus and colluvium.

Ta SEVER RIVER FORMATION (PLIOCENE AND MIOCENE)—Pebbly to moderately consolidated, gray to tan conglomerate, sandstone, and siltstone.

Tm MOUNT BELKNAP VOLCANICS (MIOCENE)—Crystalline member—Moderately welded alkali-rhyolite ash-flow tuff containing 30 percent phenocrysts of quartz (3 percent), anorthoclase (24 percent), sodic plagioclase (2 percent), and biotite (4 percent). Outer facies from the Mount Belknop caldera (Cunningham and Steven, 1979). K-Ar age is 19.0 ± 1.2 m.y. (Steven and others, 1979).

Tn Mount Baddy Rhyolite Member—Crystalline member—Moderately welded alkali-rhyolite ash-flow tuff containing nearly a fine granular mosaic of quartz and alkali feldspar, and minor plagioclase, biotite, and hornblende. Intracaldera facies of the Mount Belknop caldera.

Td Volcaniclastic rocks—Dominated by small flow breccias derived from nearby lava flows of the Mount Baddy Rhyolite Member (Tm). Some landslide debris and faulted sand and gravel are mapped with unit. Intracaldera facies of the Mount Belknop caldera.

Tc Joe Lost Tuff Member—Partly welded, crystalline alkali-rhyolite ash-flow tuff containing 12 percent phenocrysts of quartz, sodic plagioclase, and biotite, and traces of hornblende. Outflow facies of the Mount Belknop caldera.

Tn Middle tuff member—Moderately welded, crystalline alkali-rhyolite ash-flow tuff containing 12 percent phenocrysts of quartz, sodic plagioclase, and biotite, and traces of hornblende. Outflow facies of the Mount Belknop caldera.

Tm Blue Lake Rhyolite Member—Crystalline member—Moderately welded, crystalline alkali-rhyolite ash-flow tuff lithologically similar to those in the Mount Baddy Rhyolite Member (Tm). Contorted flow layers are common. Intracaldera facies of the Mount Belknop caldera.

Tl Lower tuff member—Moderately welded, crystalline alkali-rhyolite ash-flow tuff lithologically similar to the Joe Lost Tuff Member (Tm). Intracaldera facies of the Mount Belknop caldera.

Tc CONGLOMERATE (MIOCENE)—Conglomerate and sandstone. Most rounded clasts were derived from the Bullion Canyon Volcanics. Locally marks the location of the reexposed topographic wall of the Big John caldera.

Td RHVOLITE Dikes (MIOCENE)—White to light-yellow rhyolite containing phenocrysts of quartz with beta morphology, partially welded sandstone, and minor sodic plagioclase in a matrix of altered sandstone and quartz. K-Ar age is 21.9 ± 1.2 m.y. (Cunningham and others, 1978).

Td RHVOLITE Dikes (MIOCENE)—Light-gray rhyolite containing phenocrysts of Carlsbad-twinned sandstone, andesine, quartz with beta morphology, and chloritized hornblende in a matrix of altered sandstone, quartz, and minor accessory zircon and apatite.

Ti INTRUSIVE ROCKS (MIOCENE)—Dark gray and dark brown, equigranular to porphyritic intrusive rock, largely quartz monzonite, consisting of plagioclase, orthoclase, quartz, and apatite, with or without hornblende and biotite. Accessory minerals are zircon, apatite, and Fe-Ti oxides.

Td MOUNT DUTTON FORMATION (MIOCENE AND OLILOCENE)—Dark gray, dark brown, and black lava flows, flow breccias, and volcanic mudflow breccias. Largely pyroxene andesite containing phenocrysts of andesine and augite.

To OSRIS TUFF (MIOCENE)—Gray and reddish-brown, densely welded, crystalline ash-flow tuff. Phenocrysts consist of andesine (52 percent), hornblende (9 percent), biotite (4 percent), and Fe-Ti oxides. K-Ar age is about 22 m.y. (Fleck and others, 1975).

Tm BULLION CANYON VOLCANICS (MIOCENE AND OLILOCENE)—Heterogeneous lava flows and breccias ranging from rhyolite to basaltic andesite. Lava flows and breccias contain phenocrysts of plagioclase, biotite, and ilmenite, to fine-grained dark-gray rocks of intermediate composition containing small phenocrysts of plagioclase and ilmenite.

Tu Upper member—Mostly rhyolite to andesite lava flows and local ash-flow tuffs.

Tl Red lineate tuff member—Thin red lineate, densely welded ash-flow tuff that overlies the Delano Peak Tuff Member (Tbd) and flows down the topographic wall of the Big John caldera.

Td Delano Peak Tuff Member (MIOCENE)—Densely welded, crystalline quartz latite ash-flow tuff containing phenocrysts of andesine (52 percent), hornblende (9 percent), biotite (4 percent), and less than 1 percent each of quartz, zircon, and apatite. Outflow and intercaldera facies of the Big John caldera. K-Ar age is 21.8 ± 1.0 m.y. (Steven and others, 1979).

Tm Middle member—Mostly light-gray and brown rhyolite lava flows, flow breccias, and volcanic mudflow breccias that lie between the overlying Delano Peak Tuff Member (Tbd) and underlying Three Creeks Tuff Member (Ttd).

Ttd Three Creeks Tuff Member (Oligocene)—Densely welded, light-gray and brown, crystalline quartz latite ash-flow tuff containing phenocrysts of plagioclase (33 percent), hornblende (9 percent), biotite (3 percent), and quartz (2 percent). Fe-Ti oxide minerals, sandstone, and other accessory minerals occur in traces. K-Ar age is 27 m.y. (Steven and others, 1979).

Tl Lower member—Mostly volcanic mudflow breccias, flow breccias, and silty-sandstone rocks that occur below the Three Creeks Tuff Member (Ttd).

Tc CONGLOMERATE (OLIGOCENE TO PALEOCENE)—Light-gray to buff pebble conglomerate containing rounded clasts of andesite and limestone derived from Miocene and Paleocene rocks. Locally contains rhyolite sandstone.

Jn ARAPAHOE FORMATION (MIDDLE JURASSIC)—Light-gray limestone and shale and locally interbedded red to brown sandstone and interformational limestone conglomerate layers.

Jm CHINLE FORMATION (UPPER TRIASSIC)—Green, purple, and red sandstone, siltstone, and mudstone. Laminated and thinly bedded.

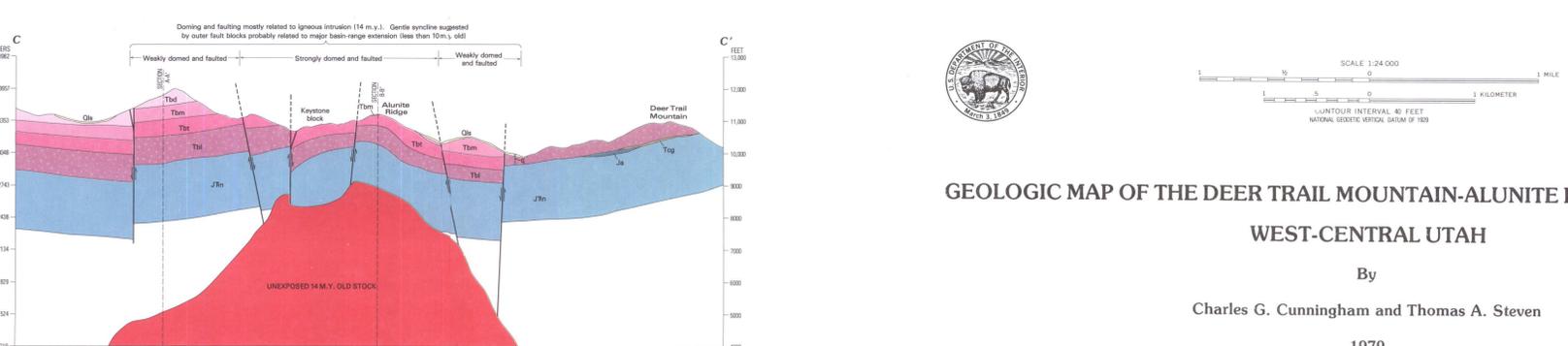
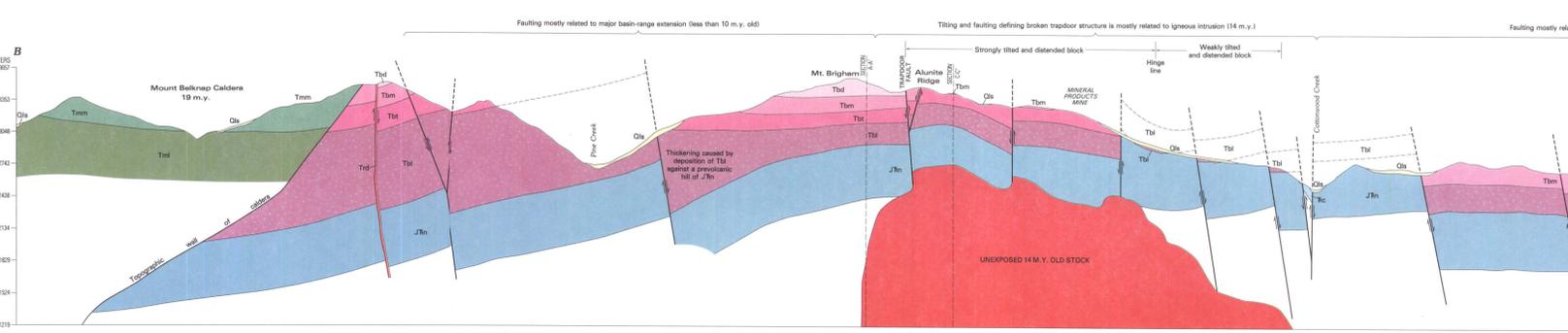
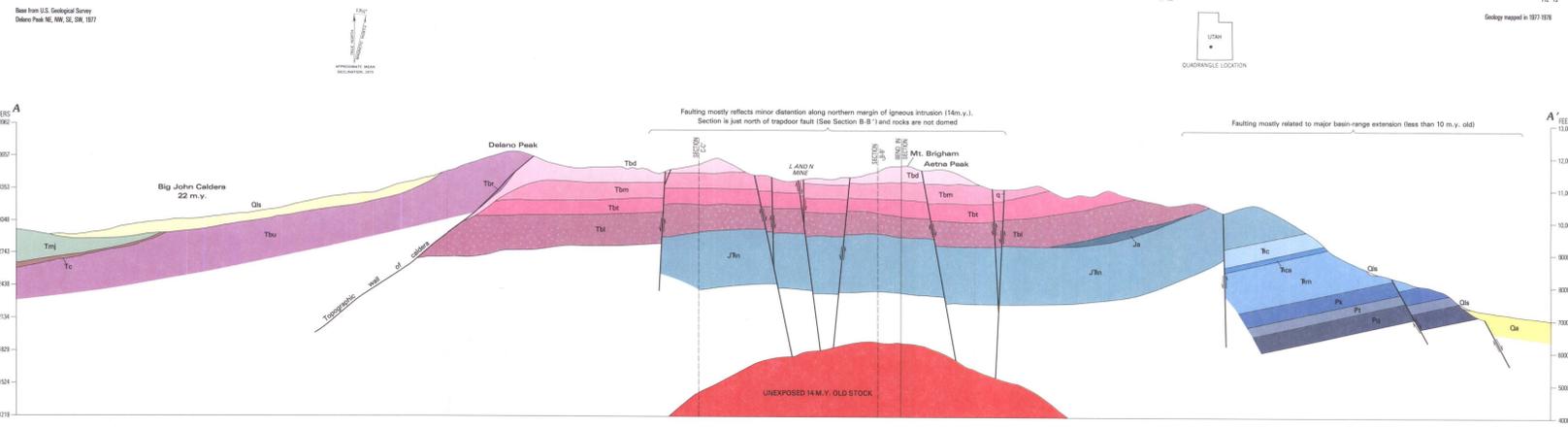
Jp Shinrump Member—Light brown to brownish-gray crossbedded poorly sorted sandstone. Crossbedding dips northwest.

Jt MOENKOPF FORMATION (MIDDLE AND LOWER TRIASSIC)—Greenish-gray and purple, fine-grained sandstone and underlying red siltstone and shale. Includes a prominent gray limestone bed near the base.

Jk KANAB LIMESTONE (LOWER PERMIAN)—Light to dark-gray limestone and dolomite.

Jb TOROWEAP FORMATION (LOWER PERMIAN)—Light-gray dolomite and yellowish-gray limestone, sandstone, and quartzite.

Jf QUANTOWEAP SANDSTONE OF McNAIR, 1951 (LOWER PERMIAN)—Light tan well-sorted, crossbedded orthoquartzite.



CONTACTS—Dashed where approximately located.

FAULTS—Dashed where approximately located; dotted where concealed. Bar and ball on downthrown side. Amount and direction of dip shown where known.

ALUNITE VEIN

QUARTZ VEIN—Amount and direction of dip shown where known.

STRIKE AND DIP OF SEDIMENTARY BEDS OR LAVA FLOWS

STRIKE AND DIP OF COMPACTION FOLIATION

TOPOGRAPHIC WALL OF CALDERA—Solid where it follows a contact, broken where concealed; hachures are grouped.

BREAKAWAY SCARP AT HEAD OF A COHERENT BLOCK OF SLUMPED BEDROCK

REFERENCES

Cunningham, C. G., and Steven, T. A., 1979, Mount Belknop and Red Hills calderas and associated rocks, Maryvale volcanic field, west-central Utah: U.S. Geological Survey Bulletin 1468, 34 p.

Cunningham, C. G., Steven, T. A., and Naefer, C. W., 1978, Preliminary structural and mineralogical analysis of the Deer Trail-Alunite Ridge mining area, Utah: U.S. Geological Survey Open-File Report 78-314.

Fleck, R. J., Anderson, J. J., and Rowley, P. D., 1975, Chronology of mid-Tertiary volcanism in High Plateaus region of Utah, in: Current geology of southwest High Plateaus of Utah: Geological Society of America Special Paper 160, p. 53-62.

McNair, A. H., 1951, Paleogeographic and stratigraphic relationships of northwestern Arizona: American Association of Petroleum Geologists Bulletin, v. 35, p. 525-530.

Steven, T. A., Cunningham, C. G., Naefer, C. W., and Mehner, H. H., 1979, Revised stratigraphy and radiometric ages of volcanic rocks in the Maryvale area, west-central Utah: U.S. Geological Survey Bulletin 1469, 40 p.

**GEOLOGIC MAP OF THE DEER TRAIL MOUNTAIN-ALUNITE RIDGE MINING AREA,
WEST-CENTRAL UTAH**

By
Charles G. Cunningham and Thomas A. Steven

1979

